

Монографията „Геология на България“ се отпечатва повече от 40 години след последните обобщения върху тази тема. За това време е натрупан огромен материал, частично публикуван (често в труднодостъпни издания) у нас и в чужбина. В настоящия том II са разгледани различни аспекти (стратиграфия, седиментология, магматизъм, метаморфизъм, палеогеография, палеогеодинамика) на мезозойската геология на българските земи. Наред с гледната точка на авторите се обсъждат и нерешени, и спорни проблеми, като се очертават перспективите за бъдещите изследвания.

Монографията е предназначена за специалисти от различни области на науките за Земята, както и за студенти по геология, геофизика и география.

The book “**Geology of Bulgaria**” appears more than 40 years after the last monographs on this topic. During this time an enormous material has been collected, and has been only partially published in a number of papers in Bulgarian and foreign journals.

The present Volume II treats different aspects (stratigraphy, sedimentology, igneous and metamorphic petrology, palaeogeography, geodynamics) of the Mesozoic geology of Bulgaria. The authors are distinguished researchers on the corresponding systems and series. Together with their own experience, they discuss unsolved and/or controversial problems, and trace the outlines and prospects of future studies.

The book is directed towards specialists from different domains of the Earth sciences, and students of geology, geophysics and geography.

Книгата Геология на България Том II се отпечатва със съдействието на Българската академия на науките, Геологическия институт при БАН, Министерството на околната среда и водите, Националната геоложка служба и Предприятието за управление на дейностите по охрана на околната среда.

The book **Geology of Bulgaria** Volume II is published with the assistance of the Bulgarian Academy of Sciences, the Geological Institute, the Ministry of Environment and Water, the National Geological Survey, and the Enterprise for Management of Environmental Protection Activities.



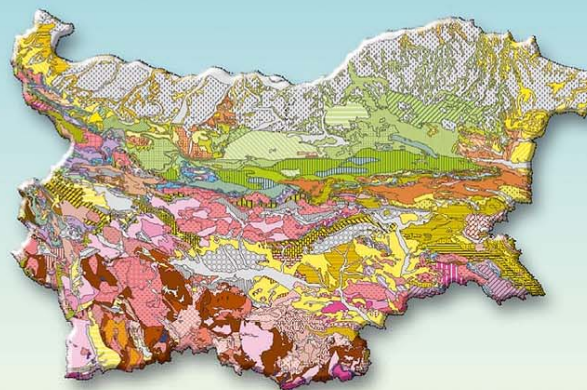
II
ТОМ

Геология на България

Геология на България

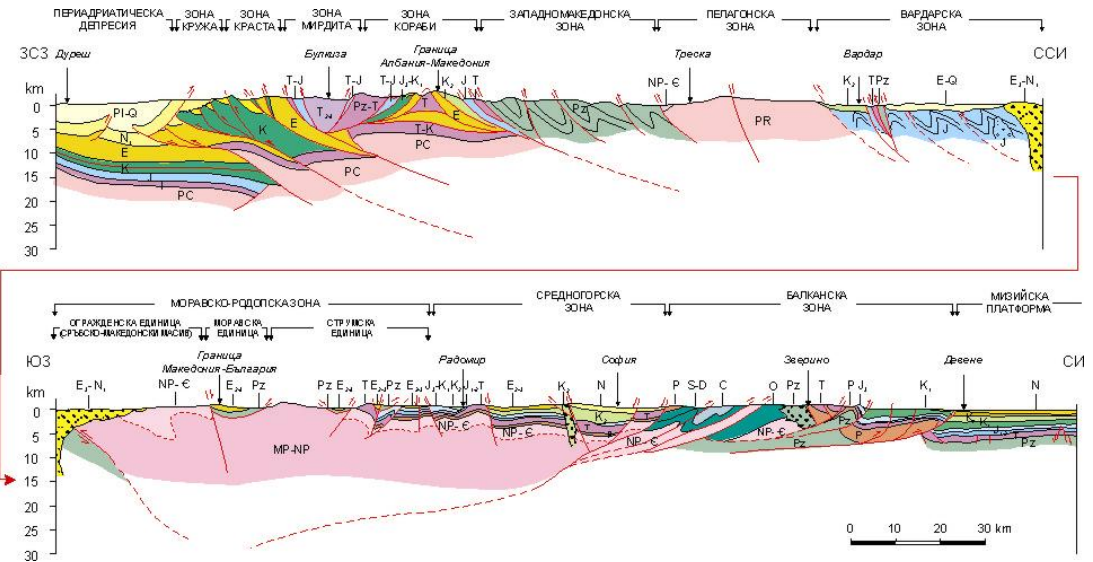
Том II

Част 5. Мезозойска геология



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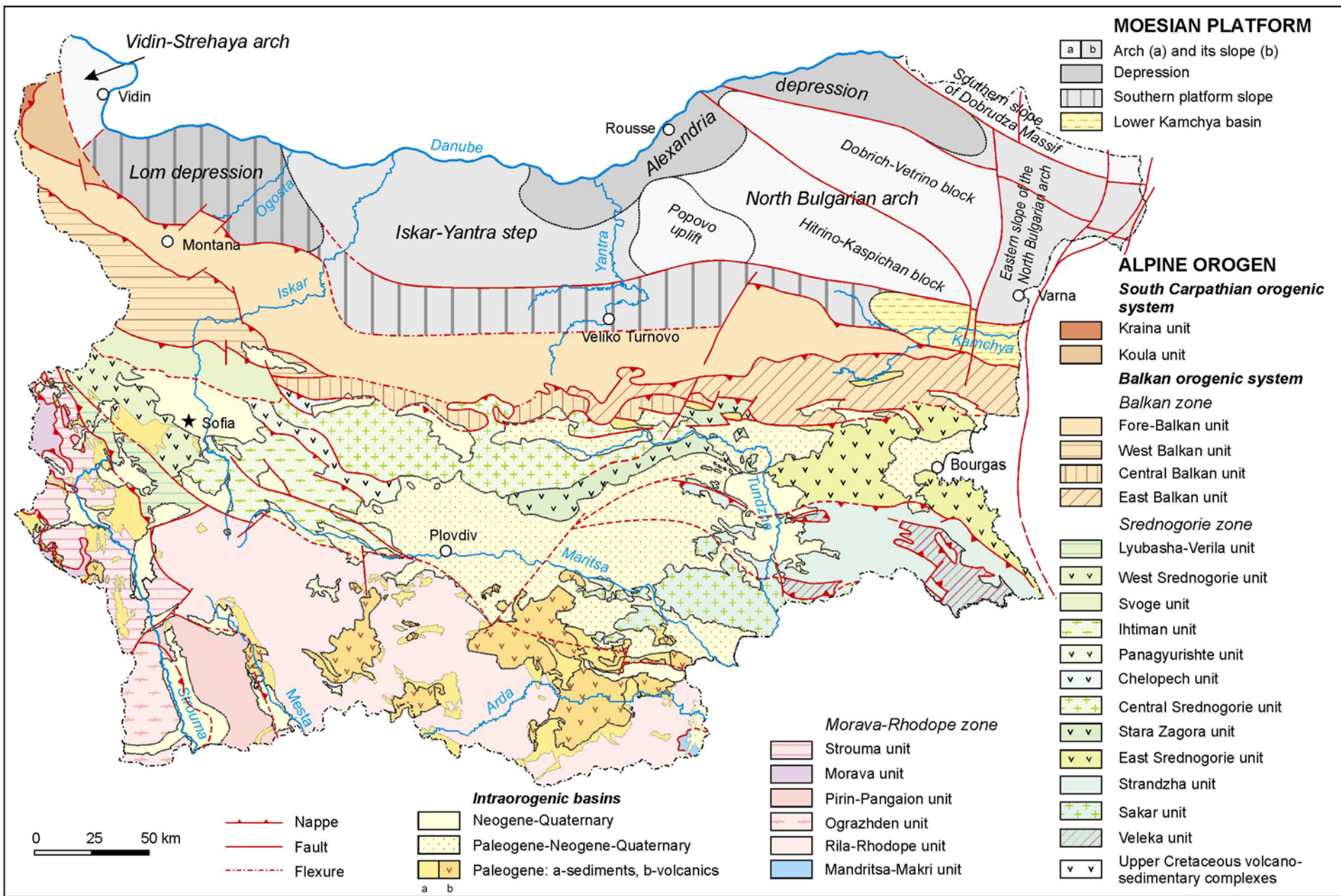
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Tectonic map of Bulgaria



The territory of Bulgaria covers parts of two major tectonic units – the **Moesian platform** and the **Alpine orogenic belt**. The structure of the platform is dominated by vast positive (arches) and negative (depressions) structures that are faulted into blocks, horsts and grabens of different rank. The Alpine orogen is a generally WNW-ESE trending belt of predominantly north-verging folds and thrust sheets. Two orogenic systems can be distinguished – Carpathian and Balkan system. They are further subdivided into units of lower rank (zones, units, subunits).

The **Moesian platform** forms the foreland of the Alpine orogen. The sedimentary cover consists of subhorizontal Mesozoic and Cenozoic, mainly shallow-marine terrigenous-carbonate and carbonate platform successions up to 7–8 km thick, which overlie with angular unconformity a gently folded Palaeozoic basement. Major unconformities at the base of the Triassic, Jurassic, Upper Cretaceous and Middle Eocene correlate well with the main compressional events within the Alpine fold-and-thrust belt.

The Alpine orogenic belt consists of dominantly north-verging thrust sheets and fold structures of different rank that resulted from multiphase collisions and related compressional events in Late Triassic, mid-Jurassic, mid-Cretaceous, Late Cretaceous and mid-Eocene time. They were followed by crustal extension, collapse of the orogen and development of a system of Cenozoic intraorogenic basins.

Two orogenic systems are distinguished in the Alpine belt – South Carpathian and Balkan. They are well spatially differentiated and show some essential differences in rock composition, structure and geological evolution.

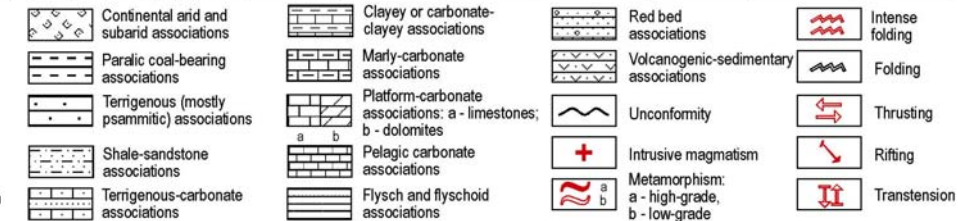
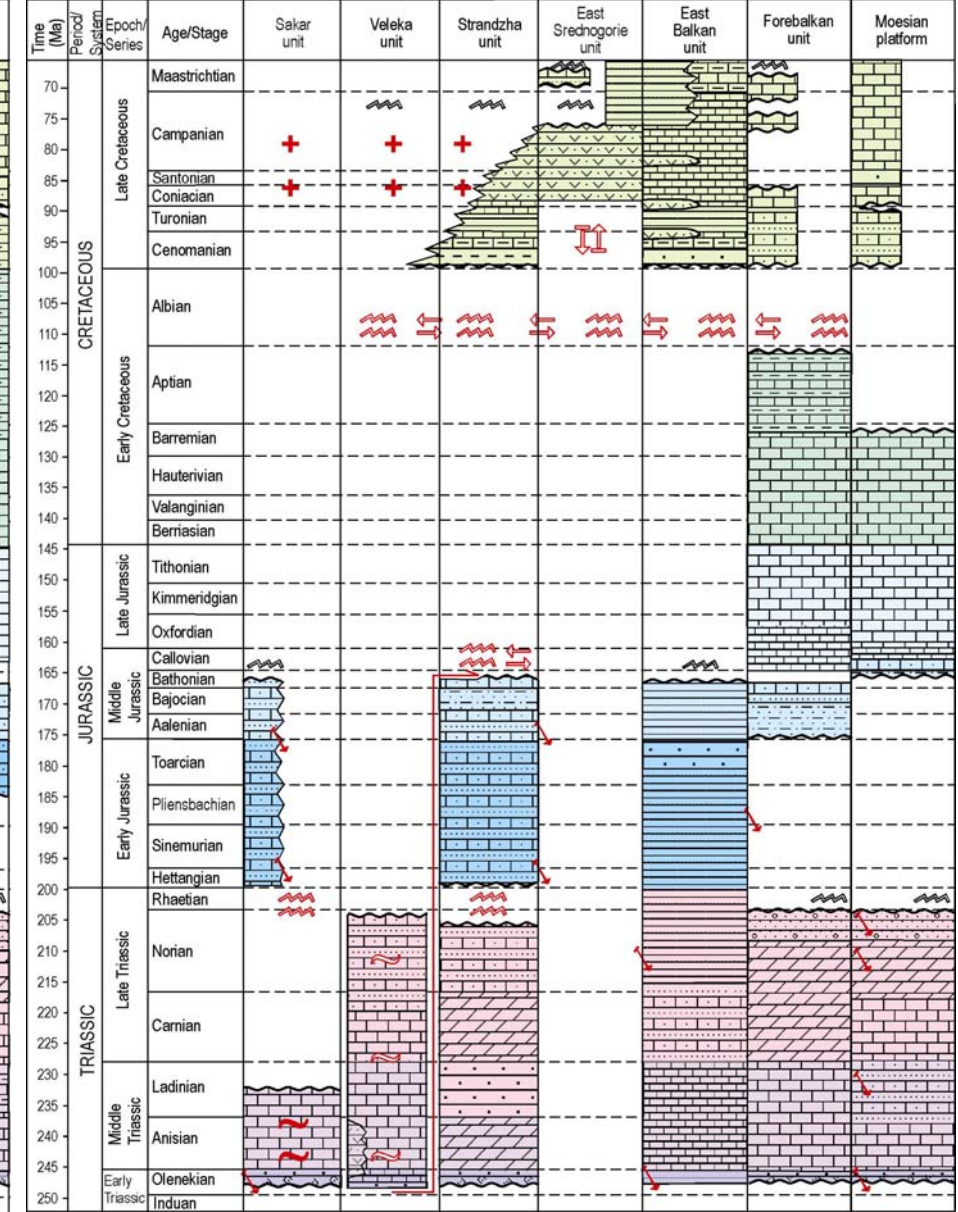
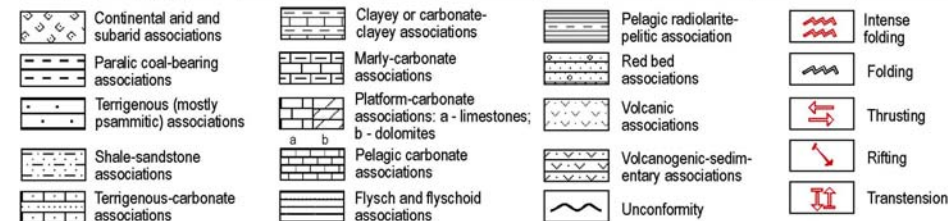
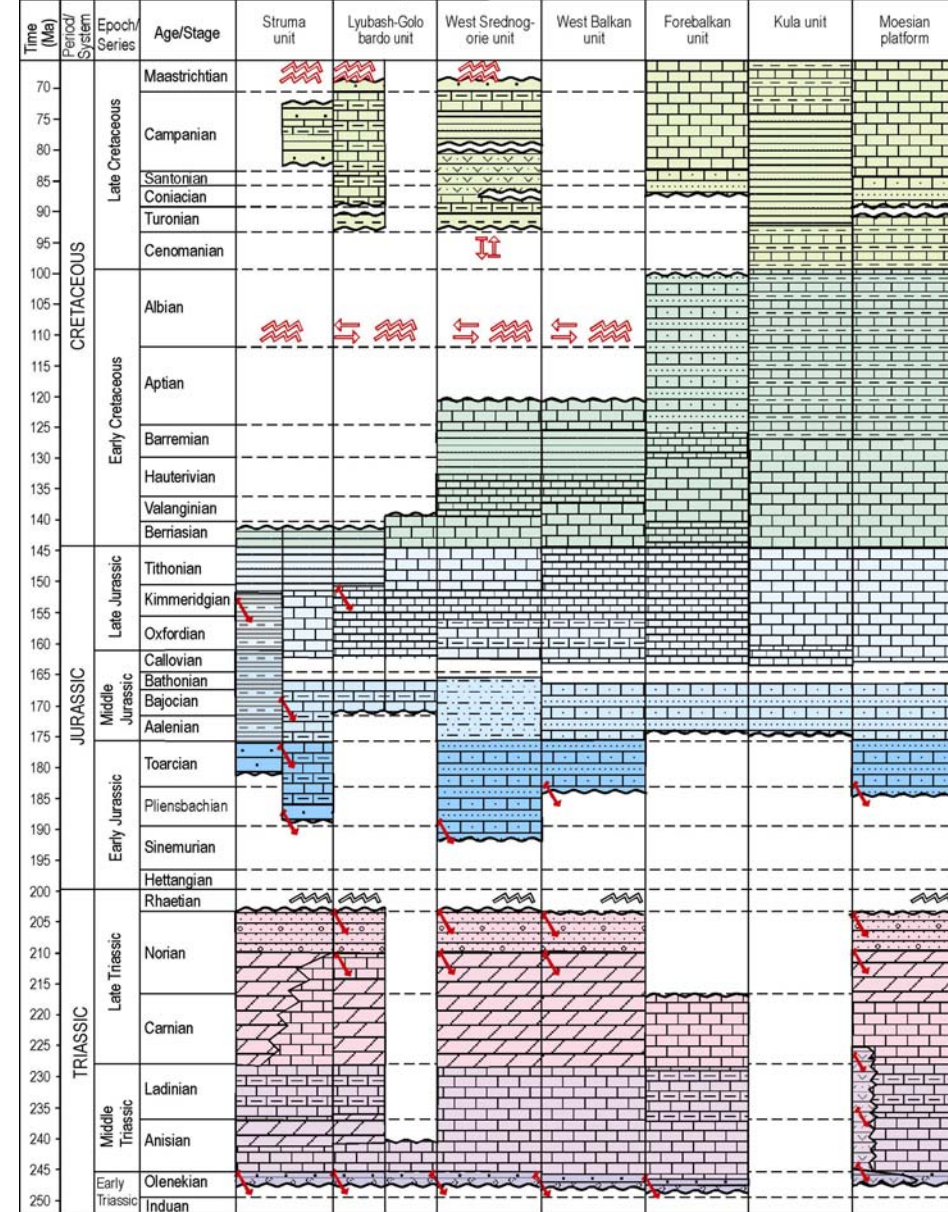
The South Carpathian orogenic system is poorly exposed only in north-westernmost Bulgaria under a thick cover of Neogene and Quaternary deposits. Two units are distinguished. *Krayna unit* is an allochthonous sheet traced as an extension of the Severin nappe in Serbia and Romania. It comprises Upper Jurassic – Lower Cretaceous flysch sequences that were deformed into east-verging folds in mid-Cretaceous time. Toward the end of the Late Cretaceous, this unit was thrust over the eastern, para-autochthonous *Kula unit*. It consists of Turonian–Senonian flysch that covers with gradual transition Albian–Cenomanian clay-carbonate sediments from the southern margin of the Moesian platform.

The Balkan orogenic system occupies the area between the Moesian platform and the units of the South Carpathians to the north, and the Vardar zone to the southwest and south. It is subdivided into three tectonic zones: Balkan, Srednogie and Morava-Rhodope.

The Balkan zone (Balkanides) forms the external northern parts of the Alpine orogen. Its typical features are: (1) wide occurrence of Triassic and Jurassic–Lower Cretaceous platform carbonates transited from the Moesian platform; (2) development of Upper Jurassic–Lower Cretaceous and Upper Cretaceous–Palaeocene flysch associations; (3) almost full absence of Mesozoic magmatic products; (4) main and final compressional events towards the end of the Middle Eocene, preceded by Late Cretaceous, mid-Cretaceous and weak Triassic deformations. The zone is subdivided into four tectonic units: Fore-Balkan, West Balkan, Central Balkan and East Balkan.

The Srednogorie zone is characterized by wide distribution of Upper Cretaceous volcano-sedimentary succession and plutonic bodies. Its northern boundary with the Balkan zone is traced by north-verging mid-Eocene reverse faults and thrusts whereas the southern one with the Morava-Rhodope zone is a system of faults (Maritsa fault zone). The main compressional events took place toward the end of the Late Cretaceous, followed by north-verging thrusting during the mid-Eocene in the northern parts of the zone. Wide areas of the Srednogorie zone are covered by Cenozoic deposits. Based on tectonic relations and specific features of the pre-Mesozoic and Mesozoic sections we distinguish 12 tectonic units.

The Morava-Rhodope zone includes fragments of six tectonic units: Morava, Ograzhden, Struma, Pirin-Pangaion, Rila-Rhodope and Mandritsa-Makri, each of them with relatively independent pre-Late Cretaceous history. The common features of these units are: (1) widely exposed high-grade metamorphic basement complexes; (2) frequent Late Cretaceous and Tertiary intrusive bodies of different size; (3) development of isolated Palaeogene basins with continental and shallow marine sediments that associate with dominantly acid and intermediate volcanic rocks; (4) main mid-Cretaceous compressional deformations followed by Late Cretaceous–Tertiary extension and exhumation; (5) thick continental crust (50–52 km) in the central parts of the zone, thinning to 34–37 km in SE and NW direction. The age of the high-grade metamorphic complexes in some units is not well constrained and controversial. Conflicting alternatives suggest Precambrian and/or Palaeozoic age of the protolith, and Precambrian, mid-Cretaceous or even Late Cretaceous – Early Tertiary age of the high-grade metamorphic events.



Main stages in the Mesozoic evolution

Based on the traditions in Bulgarian geology and the new information obtained, we divide four stages in the Mesozoic evolution of Bulgaria. They are well constrained by the major unconformities in the Mesozoic sections, and can be correlated with the global palaeogeodynamic events in the Eastern Mediterranean.

The Triassic stage began after the end of the Variscan orogeny, denudation of the orogen and formation of the Late Permian arid plain. During that stage the eastern part of the Balkan Peninsula developed as a segment of the passive European continental margin and part of the Peri-Tethyan region. During the Early Triassic Epoch braided fluvial systems of Buntsandstein type formed. Differential subsidence, accompanied by rifting, led to marine transgression towards the end of the Early Triassic, to a stable marine regime during the Middle and most of the Late Triassic, and regression in latest Triassic times. Tethyan-type deposits developed only in the East Balkan (Kotel, Mator) trough, and the arcs of the future Veleka unit (Strandzha type Triassic) and Mandritsa-Makri unit (Circum-Rhodope belt).

The Early–Middle Jurassic stage is characterized by renewal of the passive margin regime, extension and differential normal faulting that led to development of a system of horsts and grabens in the confines of the Moesian platform and the Early Cimmerian orogen to the south. This initiated gradual marine transgression, extending over the whole interval from Hettangian to Bathonian time with deposition of terrigenous sediments and shallow-marine carbonates. The sedimentation continued without break only in the Kotel (Mator) trough where thick wildflysch and olistostromes were deposited, with closure of the basin in late Middle Jurassic times.

The ***Late Jurassic–Early Cretaceous stage*** records the transition from passive to convergent continental margin as a result of beginning subduction of back-arc oceanic basins beneath the Rhodope and Serbo-Macedonian “massifs”. In front of the growing mid-Cretaceous (“Austrian”) orogen a flexural foredeep developed – the Nish-Troyan flysch trough, which accumulated turbidites (during the Kimmeridgian, Tithonian and Berriasian Age) passing to molasse deposits after the Berriasian. It was bordered to the north by the Moesian epicontinental shallow-marine basin.

Three substages are distinguished in the Early Cretaceous evolution. The first, *Berriasian–Early Barremian substage* is a continuation of the stable regime and extensional tendencies from the Late Jurassic. At the beginning of the second, *Late Barremian–Aptian substage*, a considerable differentiation of the basin started due to renewed faulting activity that resulted in the development of numerous large carbonate platforms and surrounding depressions with intense terrigenous deposition. The *Albian substage* is dominated by compressional regime (“Austrian orogeny”s.l.). The gradual narrowing of the marine basins culminated in almost total emersion above sea level.

The mid-Cretaceous (“Austrian”) orogeny is the most important tectonic event in the Mesozoic history of the eastern parts of the Balkan Peninsula. It is expressed in intense folding and thrusting, inversion and full closure of the troughs, and drying out of the epicontinental basins in the Moesian platform. Fragments of the mid-Cretaceous orogenic edifice are included in almost all later tectonic units. In them, the Early- and mid-Cretaceous tectonometamorphic events are recorded by isotopic data both in the pre-Alpine (Precambrian and Palaeozoic) cores (Rhodopes, Sakar and Struma units) and in their Triassic–Jurassic cover.

The Late Cretaceous stage was the time of continuous subduction of Tethyan oceanic lithosphere (Vardar ocean according to most authors) beneath the Rhodope and Serbo-Macedonian “massifs”. This resulted in the development of a volcanic island-arc system consisting of Morava-Rhodope frontal arc (magmatic), Srednogorie intra-arc sedimentary basin with dominantly submarine volcanic activity, Balkanide rear-arc (amagmatic) and Carpathian back-arc basin.

The Moesian epicontinental basin invaded the foreland of the mid-Cretaceous orogen (present Moesian platform and Fore-Balkan) by gradual transgression from east to west. The deposition began during the Cenomanian or Turonian Age with terrigenous and terrigenous carbonate sediments (with glauconite) passing upwards into clay-carbonate and carbonate successions.

The Carpathian basin developed since the beginning of the Early Cretaceous Epoch as a flexural foredeep in the NW slope of the Moesian platform and accumulated turbidite and later clay-carbonate deposits in the interval Turonian–Middle Eocene.

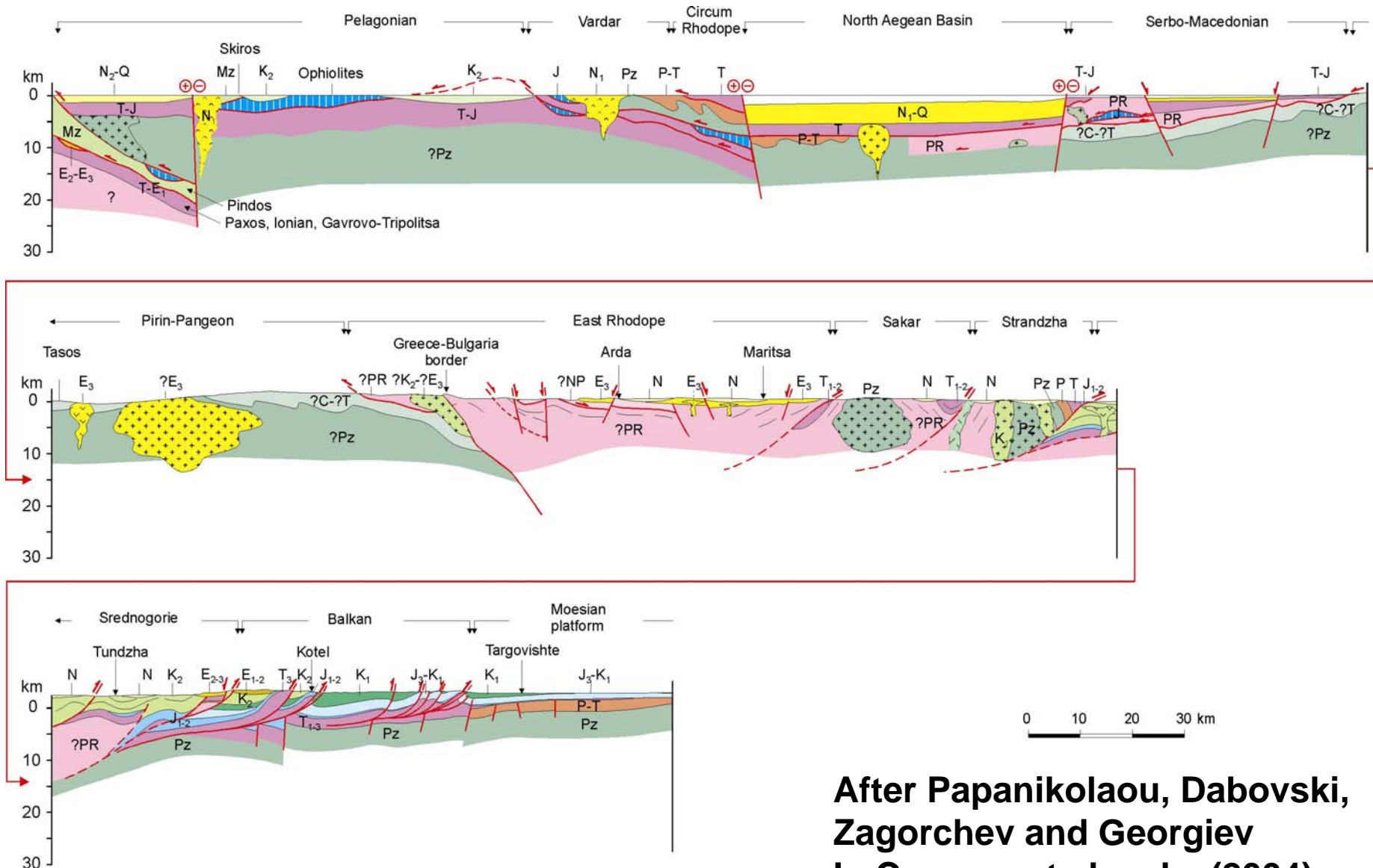
The Srednogorie basin formed and evolved upon the deeply eroded mid-Cretaceous orogen located south of the platform. Likewise the Moesian basin, the transgression advanced from east to west. Sedimentation began during the Cenomanian (East Srednogorie) or Turonian (West Srednogorie) with continental (locally coal-bearing) to shallow-marine deposits that passed upwards (Coniacian–Early Maastrichtian) into deeper-marine clay-carbonate, volcano-sedimentary and “post-volcanic” turbidite associations. The evolution of the basin terminated with regressive terrigenous deposits of Maastrichtian age in isolated “remnant” basins.

The Balkanide arc divided the Moesian and Srednogorie basins during most of the Late Cretaceous Epoch. It was a dry land, higher on the west and lower on the east, where the two basins probably merged for short periods.

The Morava-Rhodope frontal arc was also a dry land that separated the Srednogorie basin from the Vardar ocean. The high-grade metamorphic complexes of the arc were intruded by Late Cretaceous plutons of Coniacian to Maastrichtian age (by radiometric dating). There are also isolated small outcrops of Campanian volcano-sedimentary successions (East Rhodope) and Campanian sediments (SW Bulgaria).

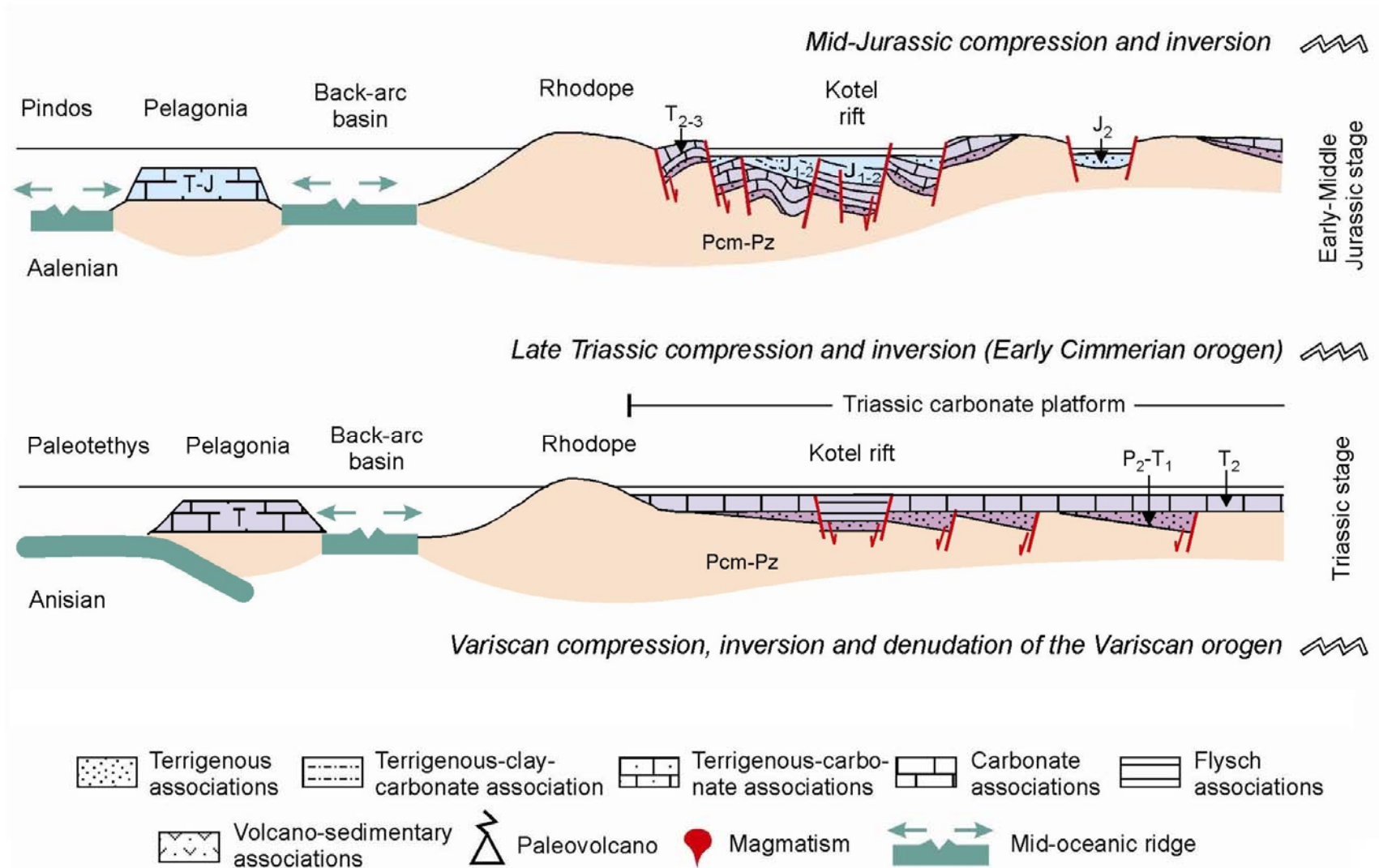
The Late Cretaceous stage, and in general the Mesozoic evolution, ended with compressive events towards the end of the Early Campanian and the end of the Maastrichtian that led to the final shaping of the Late Cretaceous orogen on the territory of South Bulgaria. In front of the orogen, the Carpathian and Emine flysch troughs developed from Late Maastrichtian to mid-Eocene time as typical foreland basins. Large parts of the orogen and the Moesian platform emerged above sea level but locally, mainly in the eastern parts of the platform and in the East Balkan, sedimentation continued without interruption also during the Paleocene.


Transect through Greece and East Bulgaria

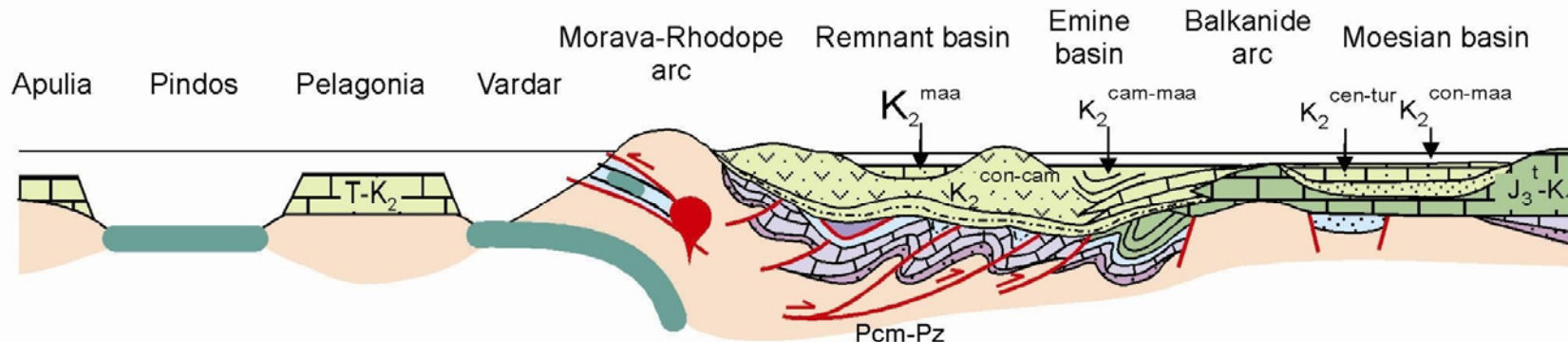


After Papanikolaou, Dabovski,
Zagorchev and Georgiev
In Cavazza et al., eds. (2004)

Schematic representation (cartoons) of the Mesozoic geodynamics in the eastern parts of the Balkan Peninsula

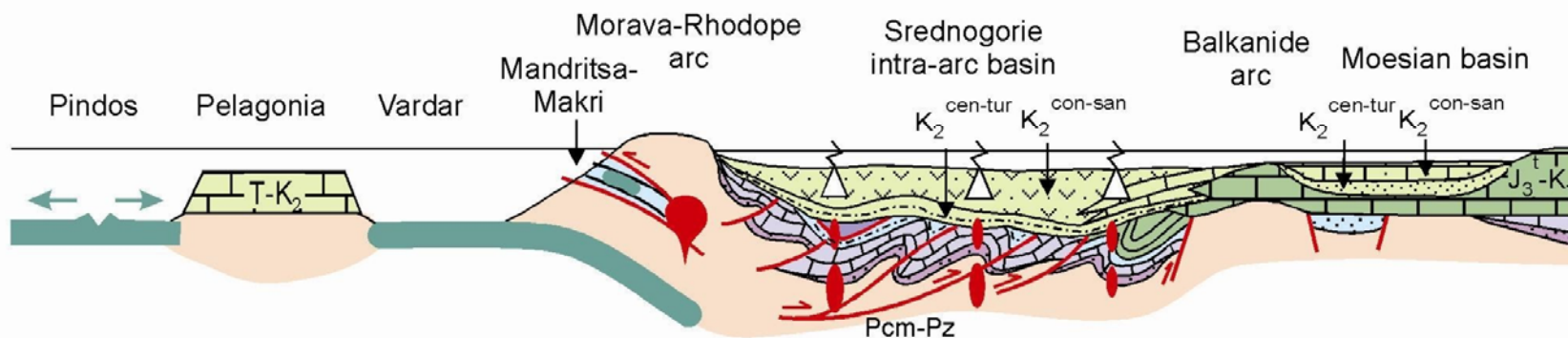



Late Cretaceous compression and inversion (Laramide orogen) 

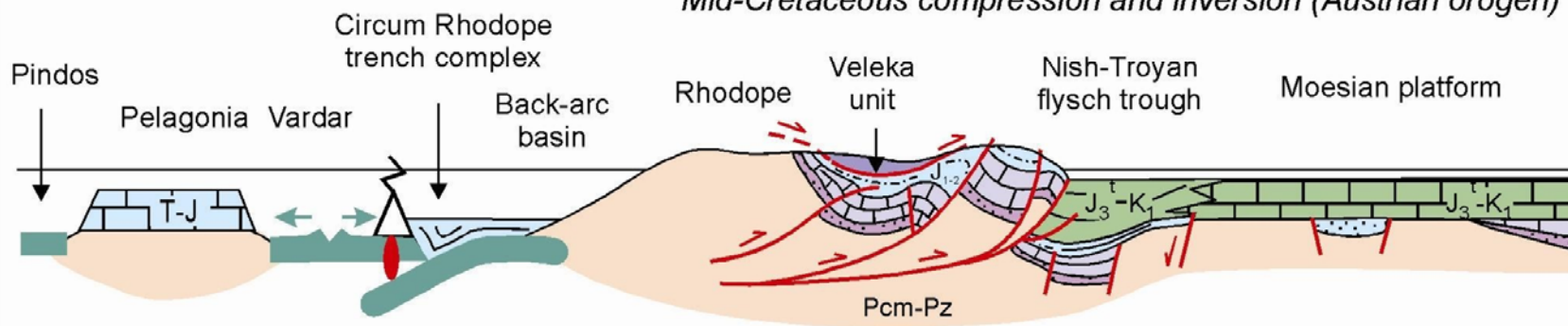


Late Cretaceous stage

Late Campanian inversion

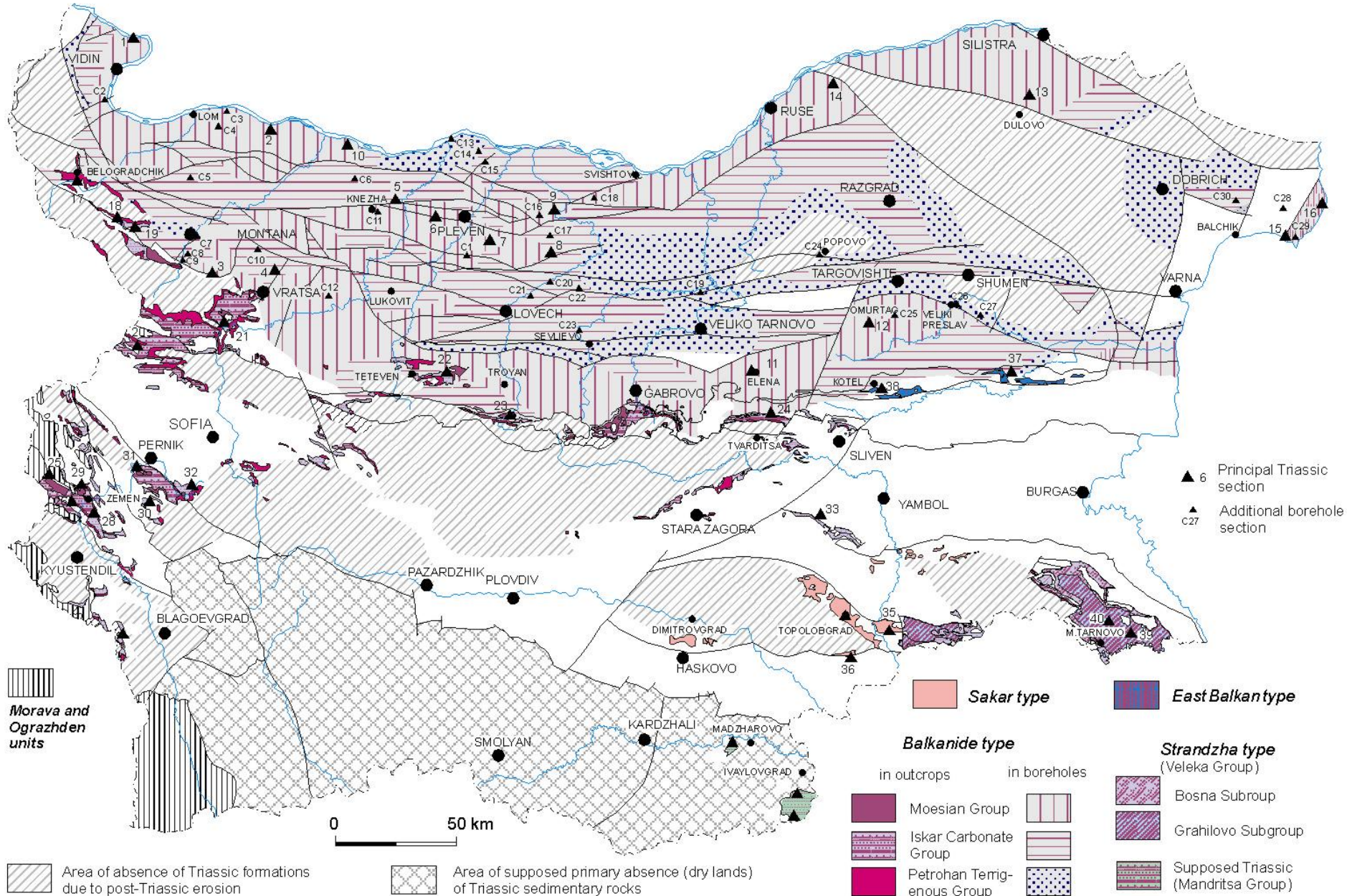


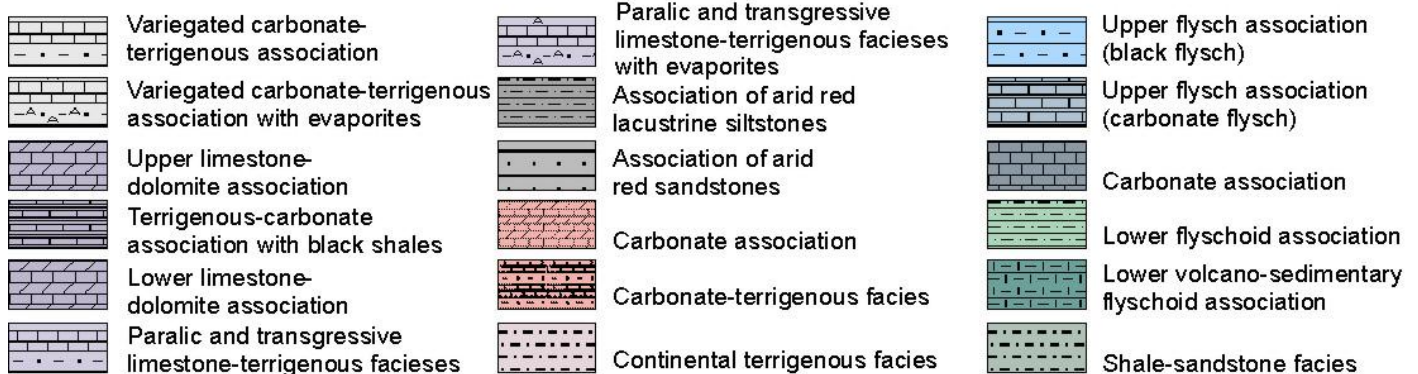
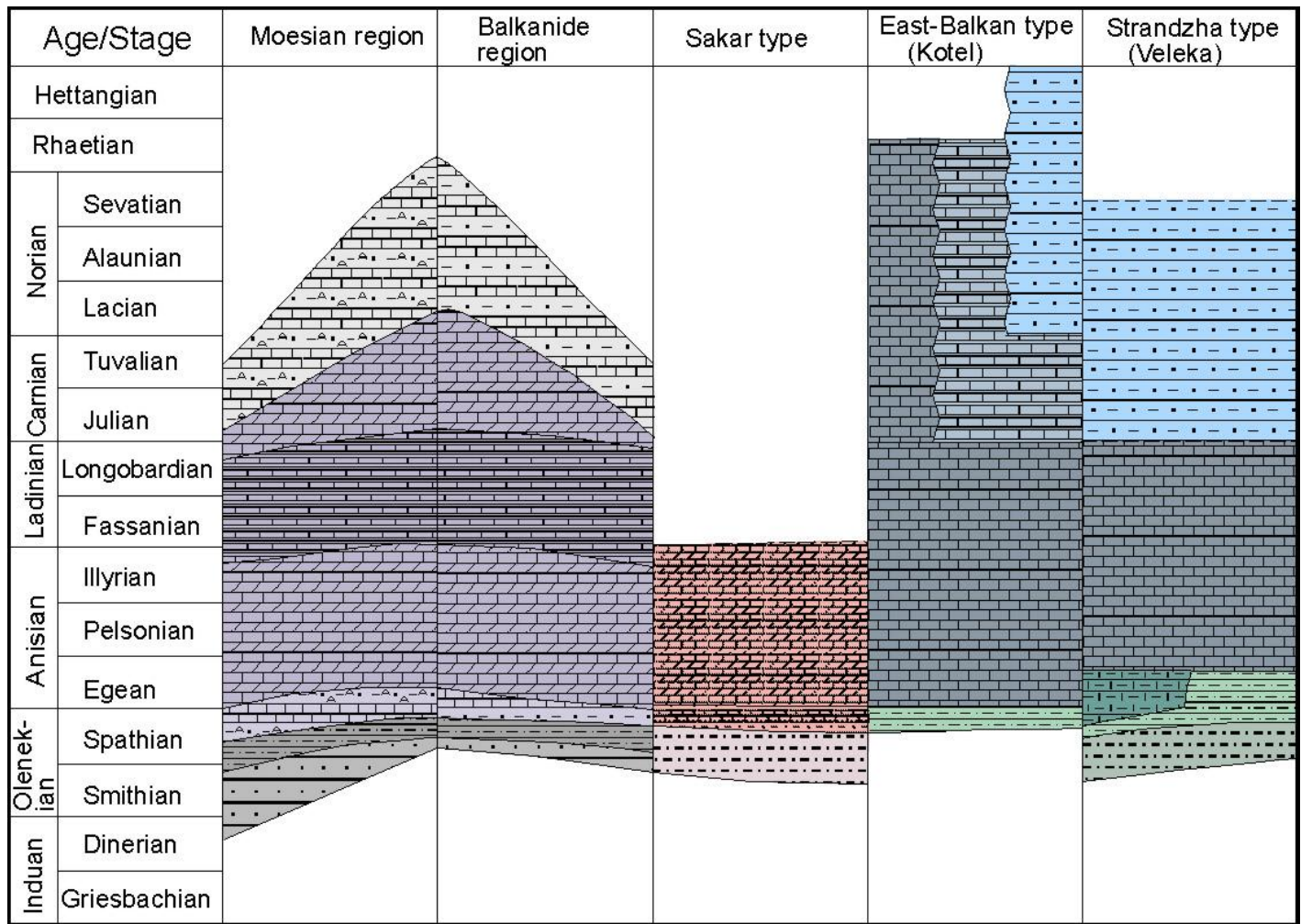
Mid-Cretaceous compression and inversion (Austrian orogen) 



Late Jurassic-Early Cretaceous stage

Occurrence of the Triassic in Bulgaria

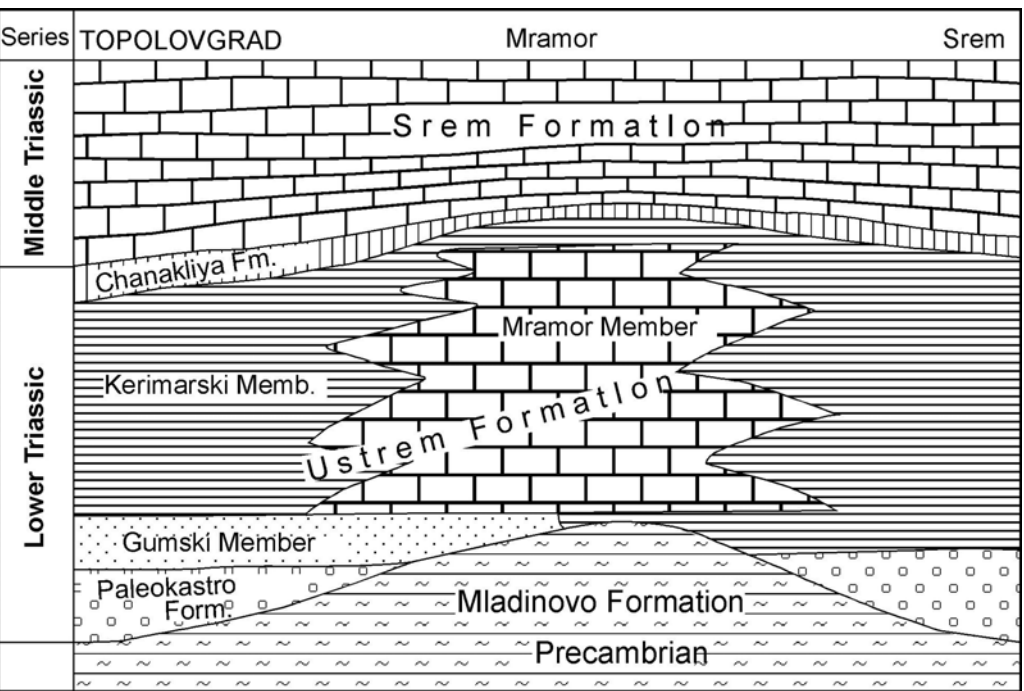




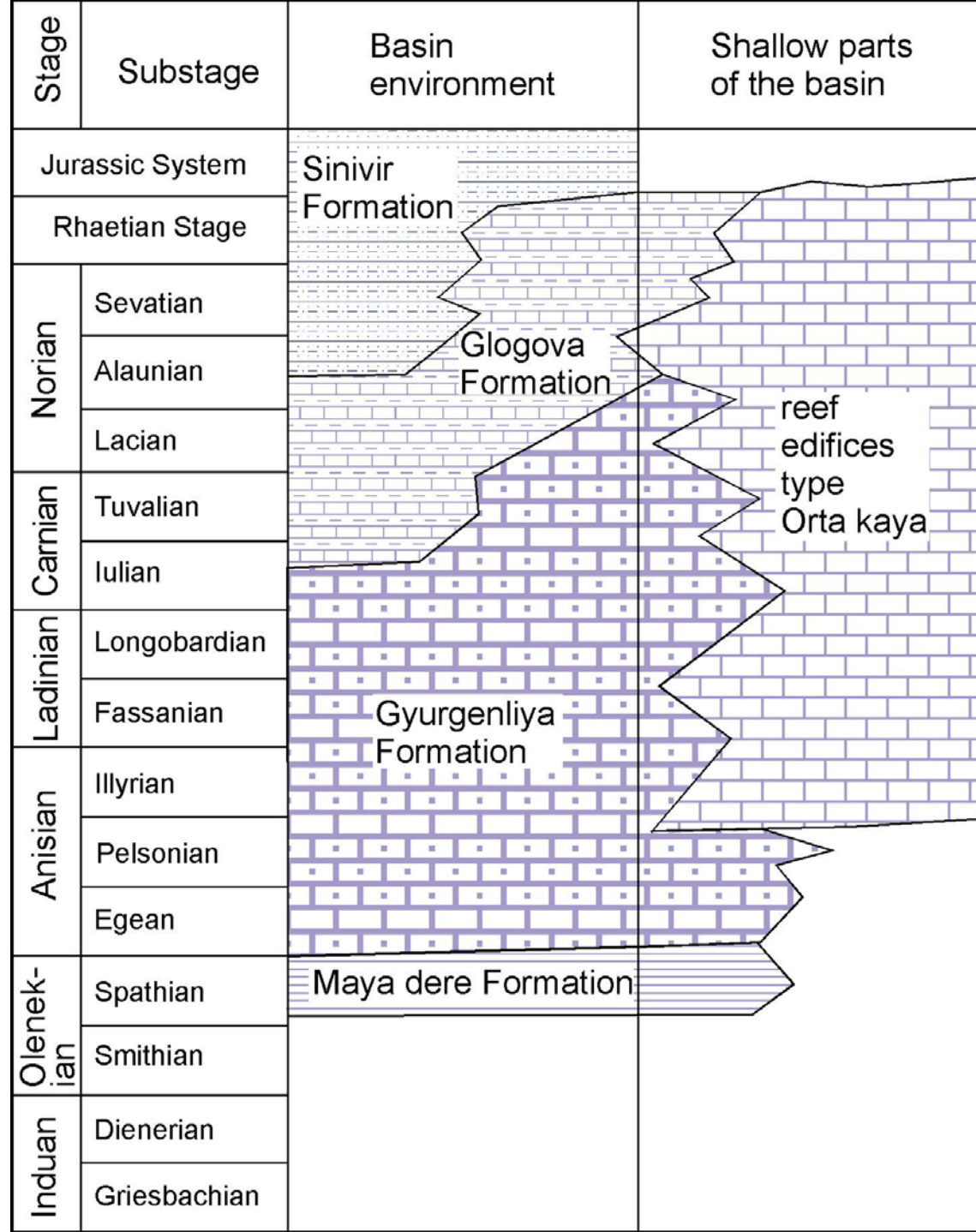
↑ **Previous page:** facies associations of the Bulgarian Triassic

Litho- and chronostratigraphy of the Strandzha-type Triassic (Veleka unit)

Stage	Substage	Чаталов (1990)	Savov et al. (1995)	Василев (1998, 2001)	Николов и др. (2009)
JURASSIC SYSTEM					
Rhaetian Stage					
Norian	Sevastian			?	
	Alaunian	Lipachka Fm.			Malko Tarnovo Fm.
	Lacian				
Camian	Tuvalian	Grafitovo Fm.	Malko Tarnovo Fm.	Lipachka Fm.	Lipachka Fm.
	Iulian			Yazminska Formation	Кондолов.
Ladinian	Longobardian	Kondolovo Fm.	Kondolovo Fm.	Kalina choukar Fm.	Lipachka Fm.
	Fassanian			Tolpanska Fm.	Lipachka Fm.
Anisian	Pelsonian				
	Illyrian			Gramatikovo Fm.	Gramatikovo Fm.
	Egean	Gramatikovo Formation	Gerovska	Gram. Mb. Zaberska Stravnitsa Fm. Kaz. Mb.	Gramatikovo Fm.
Olenekian	Spathian			Kushlyovska Fm.	Gerovska
	Smithian	Zaberska Formation	Stoilovo Formation	Zaberska Fm.	Mladezhka Fm.
Induan	Dinerian				Stravnitsa Fm.
	Griesbachian	Stravnitsa Fm.			

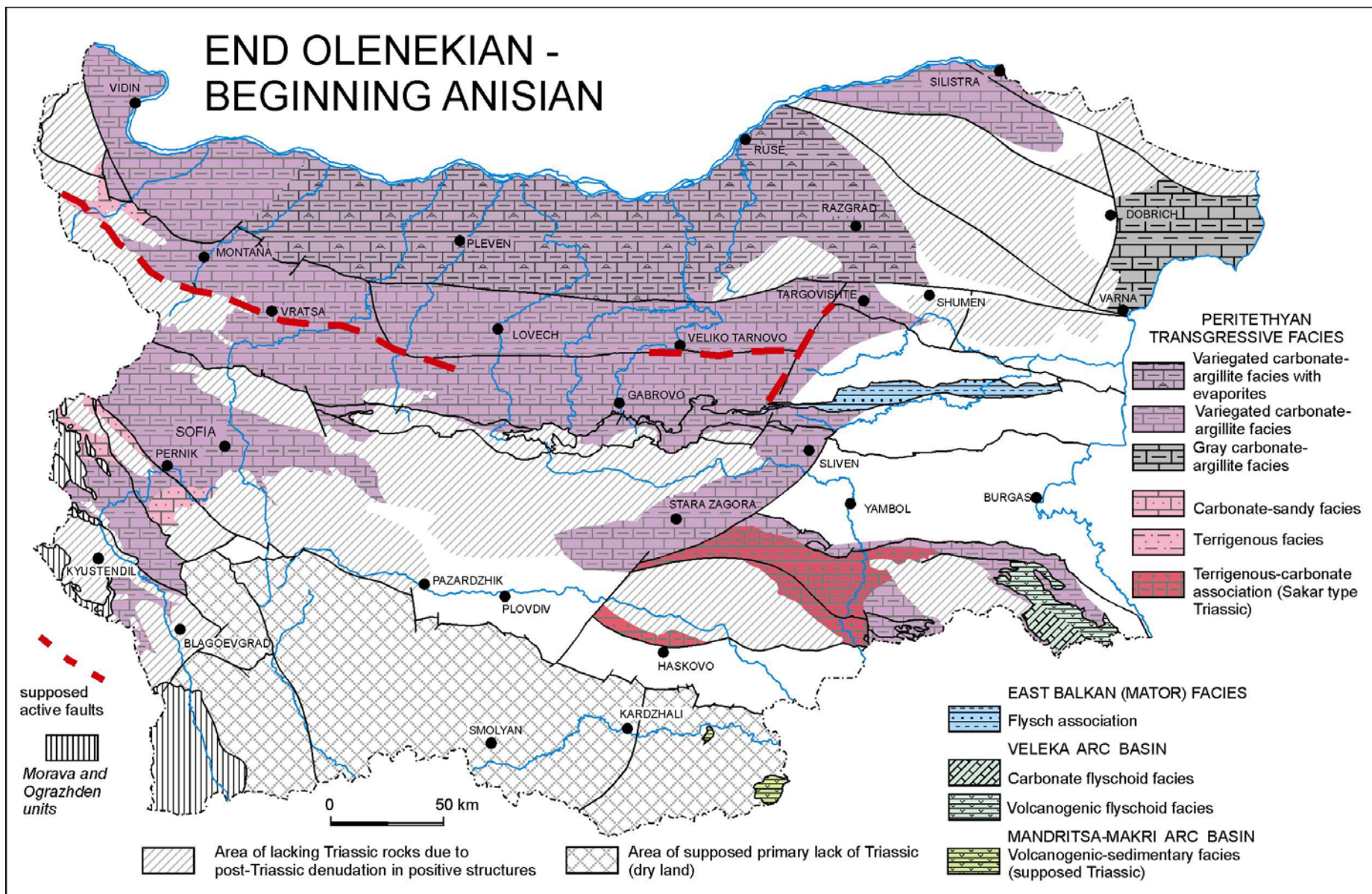


Litho- and chronostratigraphy of the Sakar-type Triassic (after Chatalov)

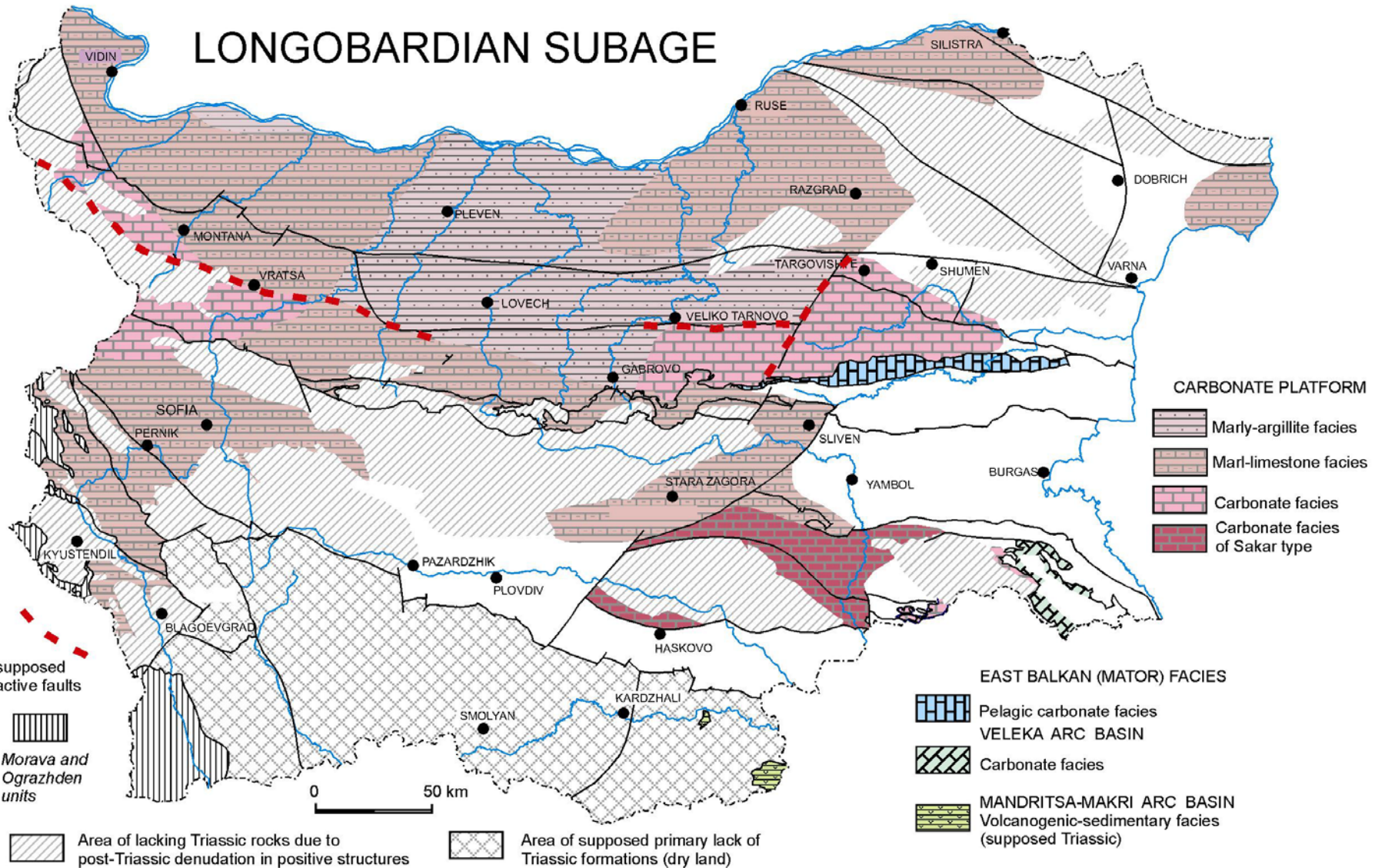


Stratigraphy of the East-Balkan Triassic type in the East-Balkan (Kotel, Mator) trough

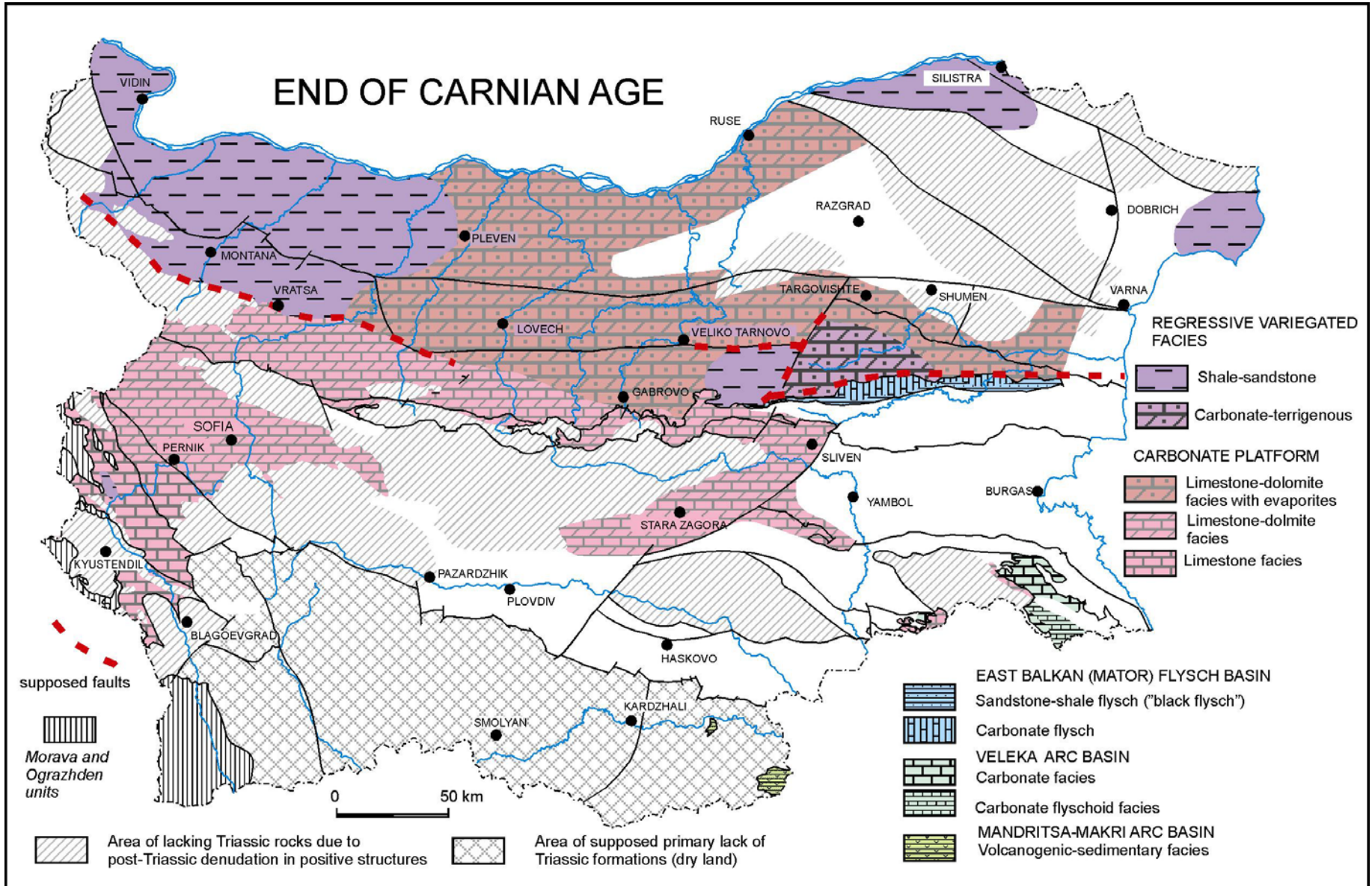
Palaeogeographic maps (modern distribution of the environments, without palinospastics)



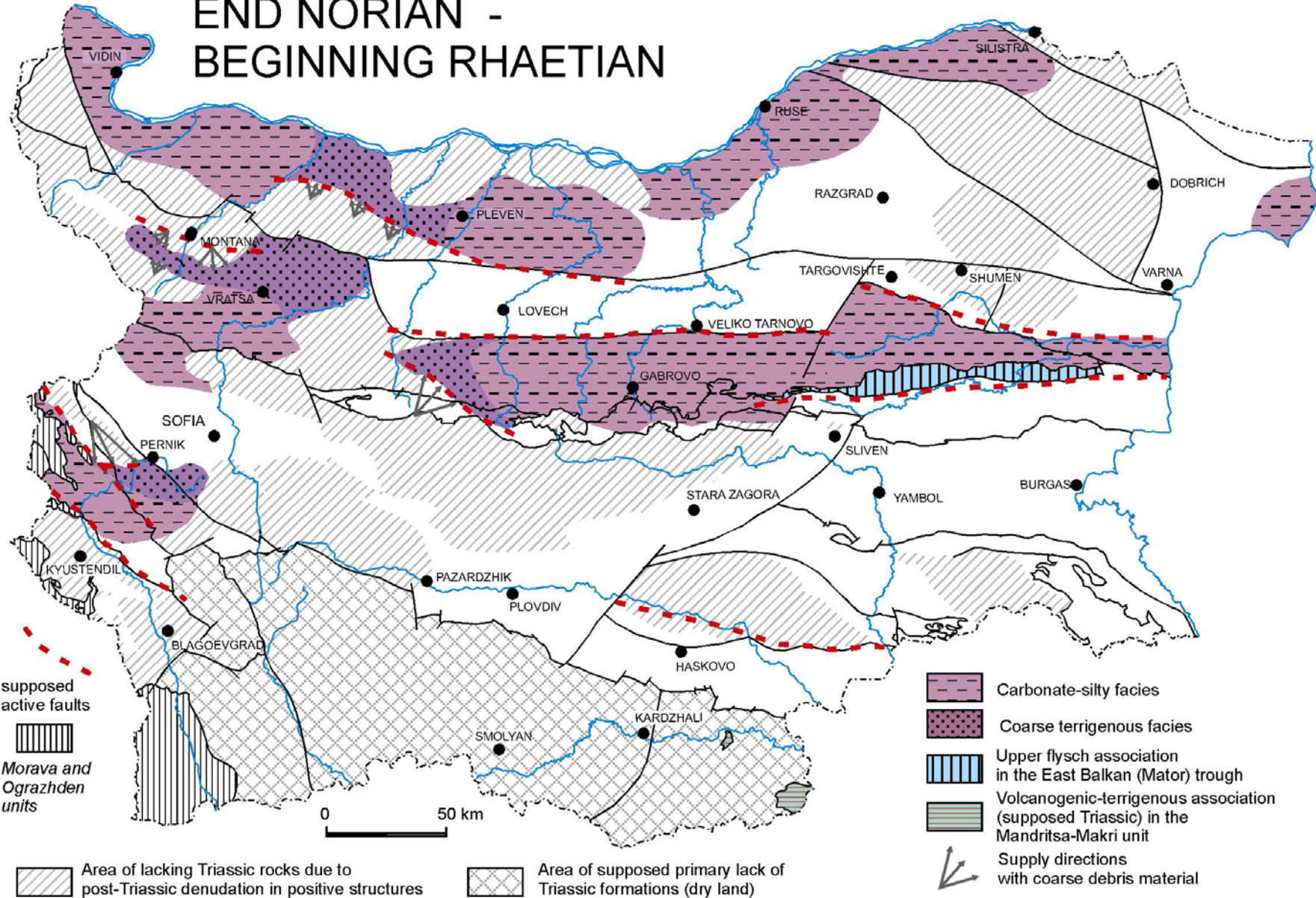
LONGOBARDIAN SUBAGE



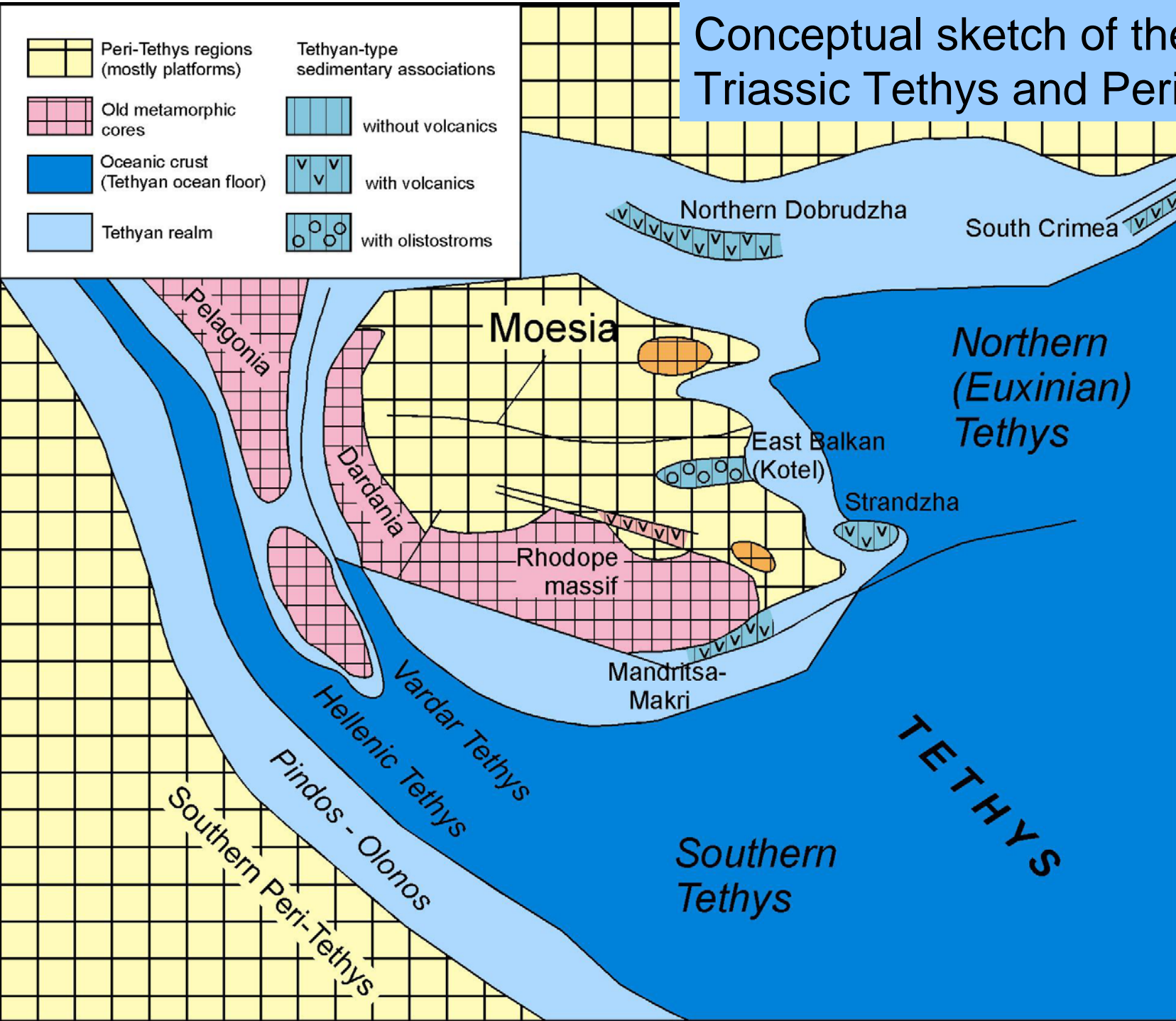
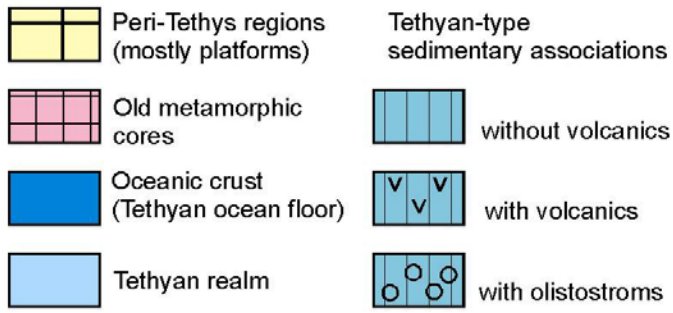
END OF CARNIAN AGE



END NORIAN - BEGINNING RHAETIAN



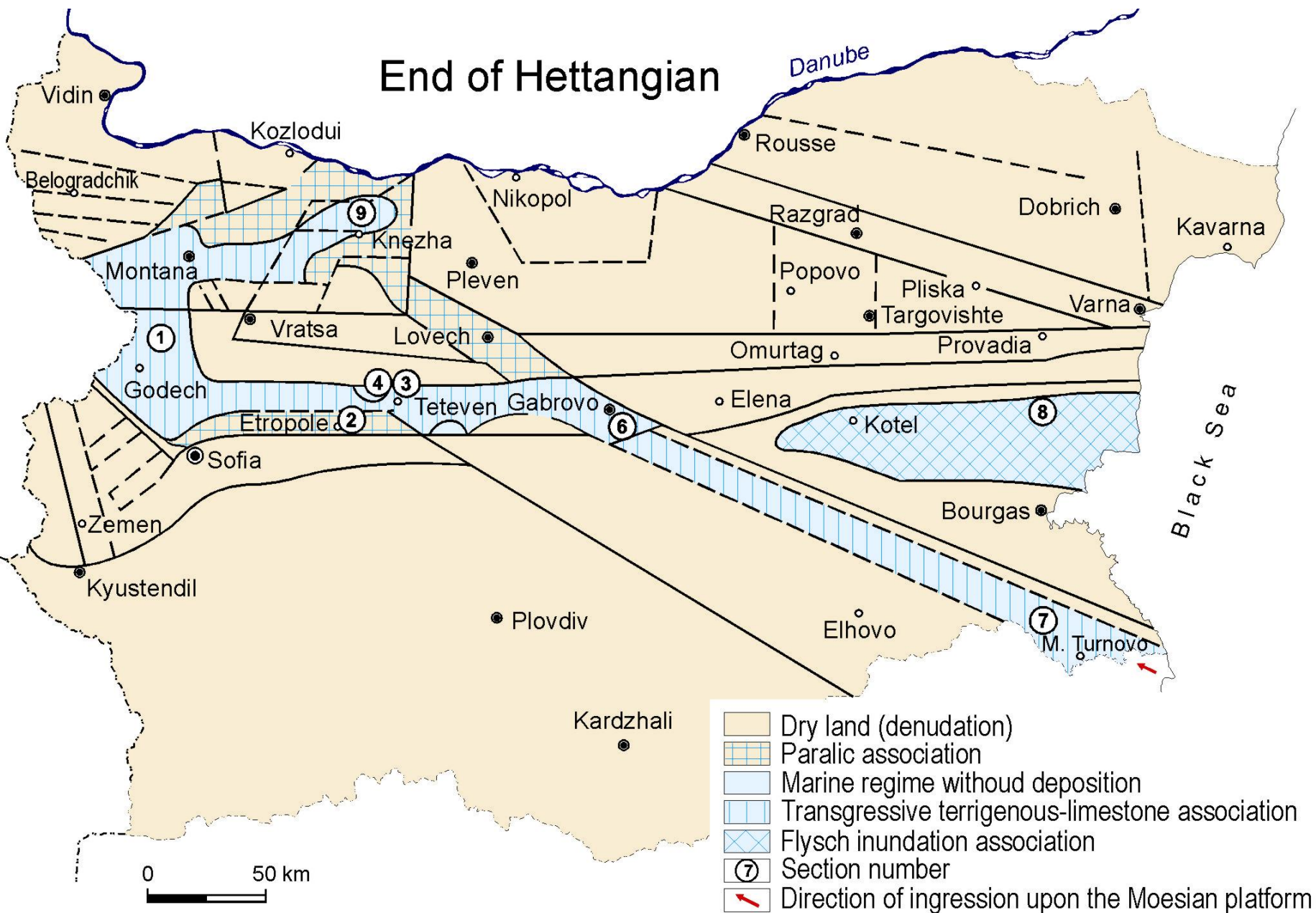
Conceptual sketch of the Triassic Tethys and Peri-Tethys



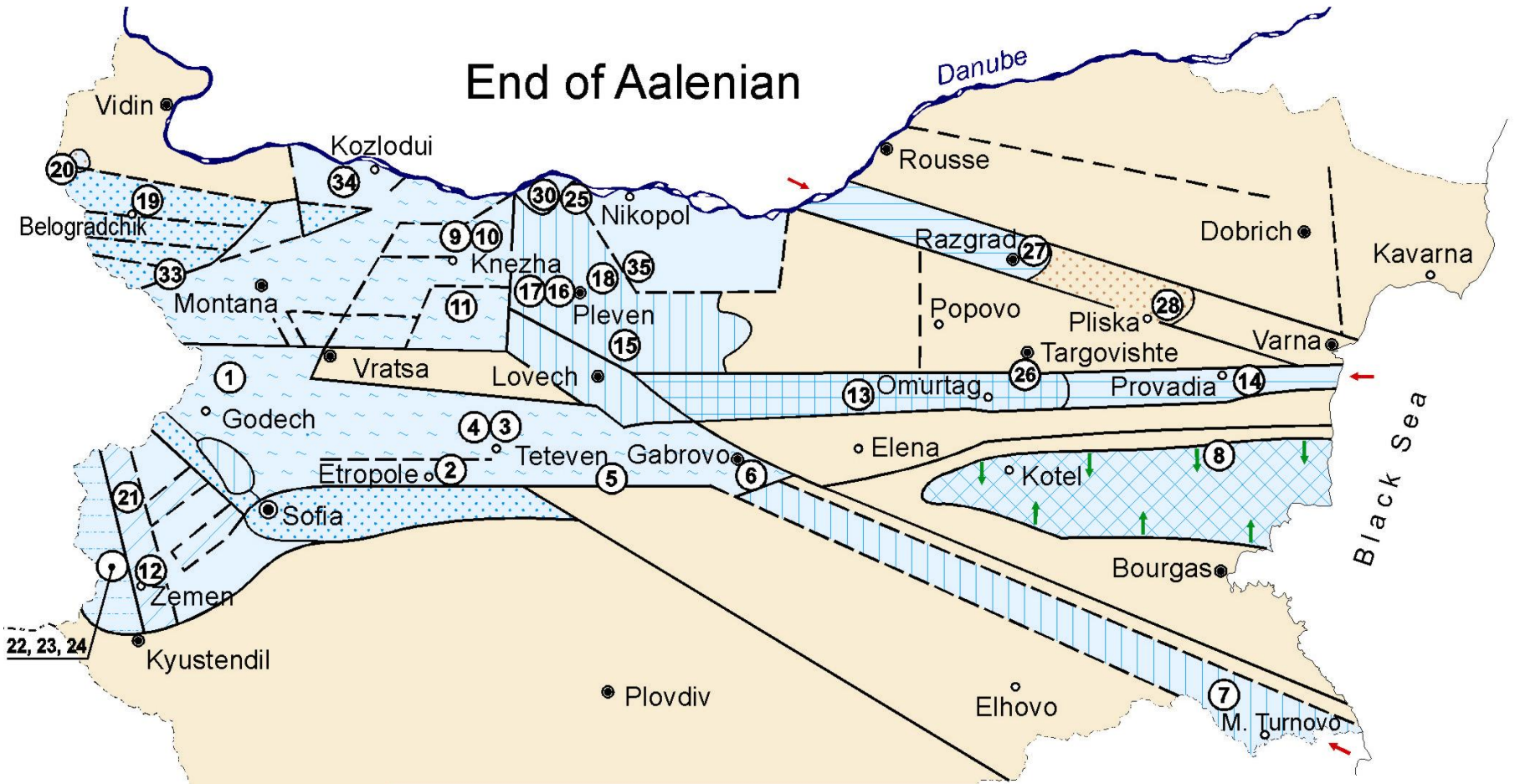
Distribution of the Jurassic in Bulgaria



Jurassic palaeogeographic maps (without palinospastics)



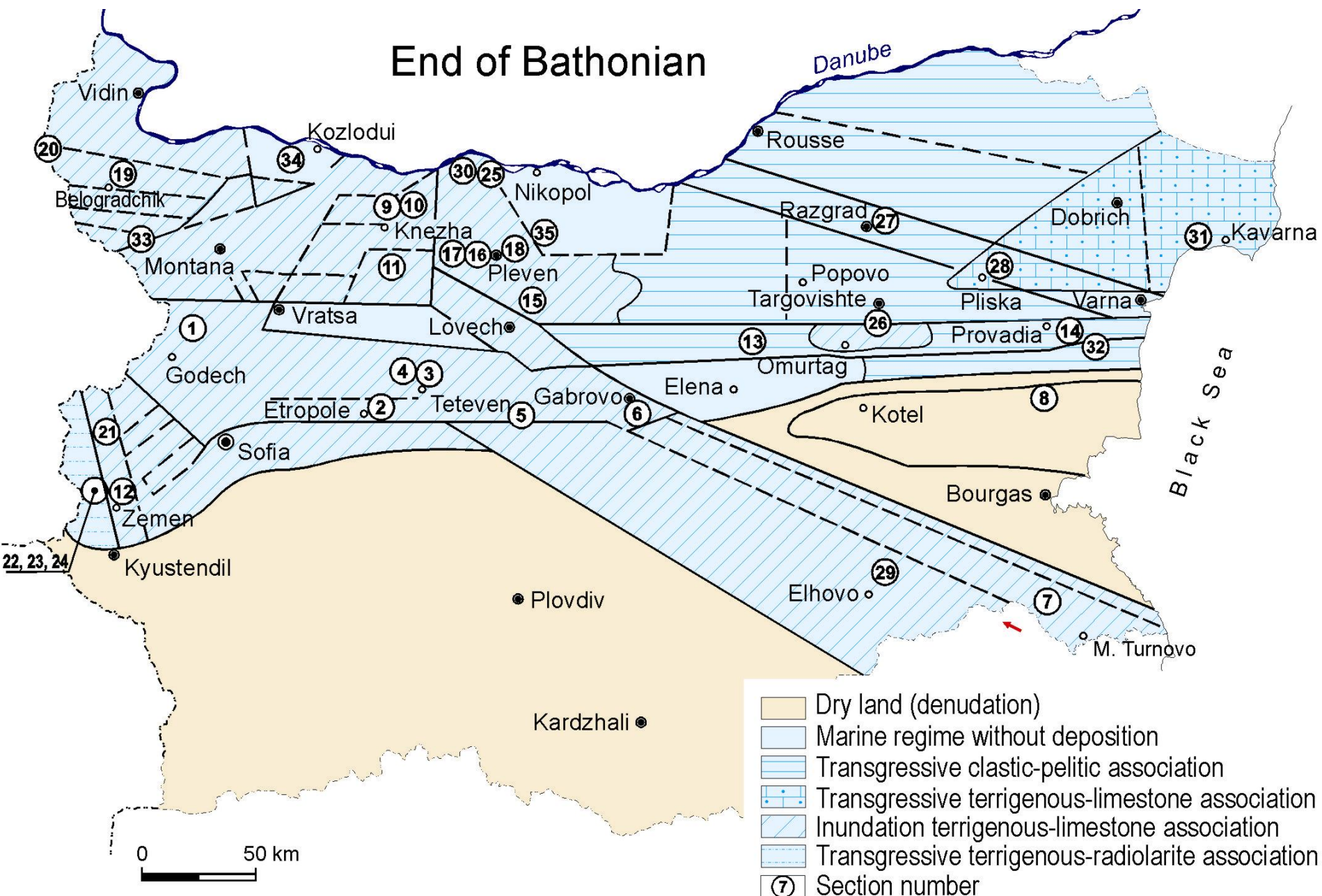
End of Aalenian



- | | | |
|---|--|---|
| Dry land (denudation) | Continental association | Paralic association |
| Marine regime without deposition | Transgressive terrigenous-limestone association | Transgressive clastic pelitic association |
| Transgressive clastic pelitic association | Intertonguing of the transgressive terrigenous-limestone and clastic-pelitic association | Inundation flysch association |
| Inundation terrigenous-limestone association | Inundation pelitic-sandstone association | Transgressive terrigenous-radiolarite association |
| Inundation pelitic-sandstone association | Transgressive terrigenous-radiolarite association | Section number |
| Transgressive terrigenous-radiolarite association | Closure of Kotel embriental rift | Direction of ingress upon the Moesian platform |



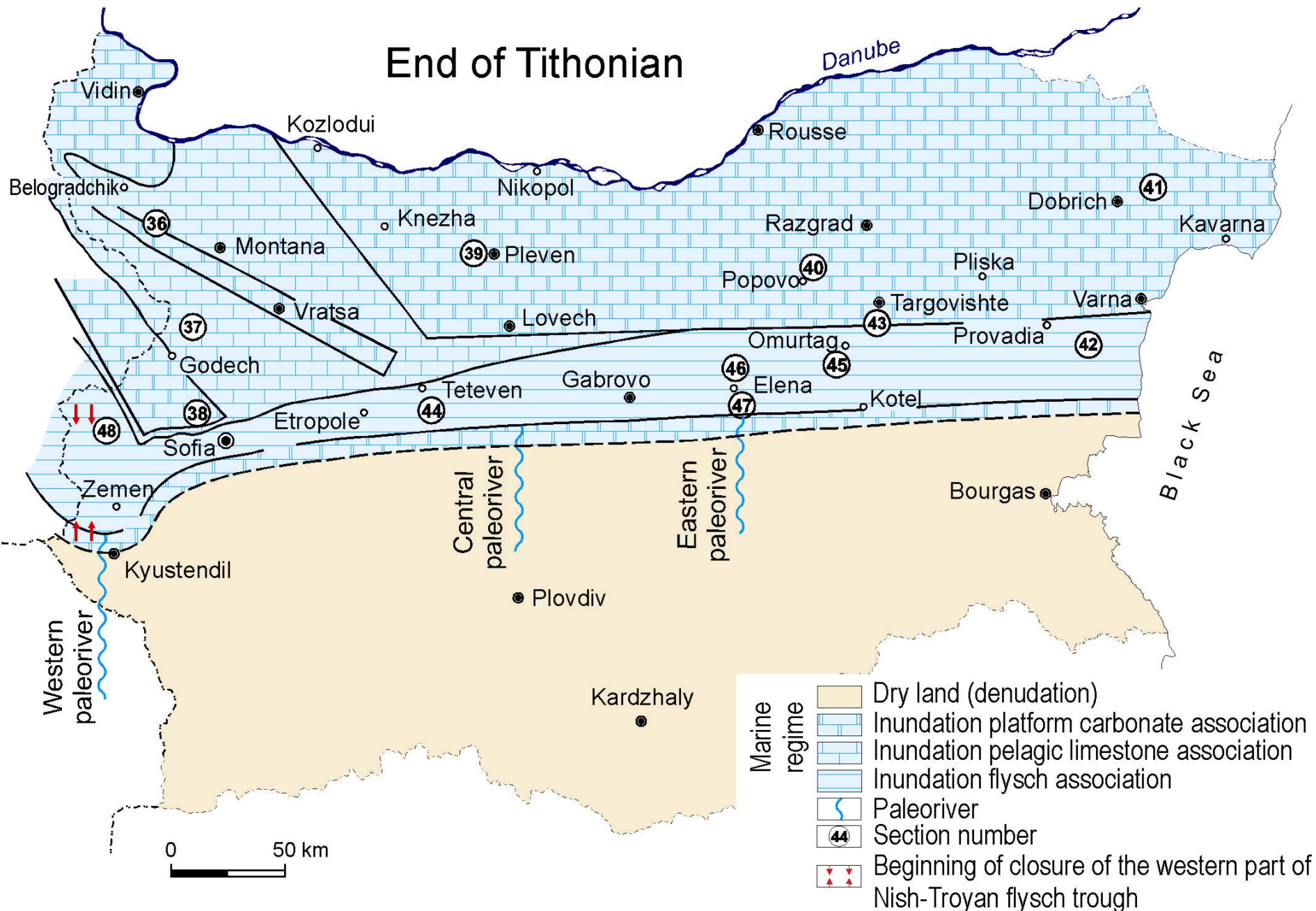
End of Bathonian



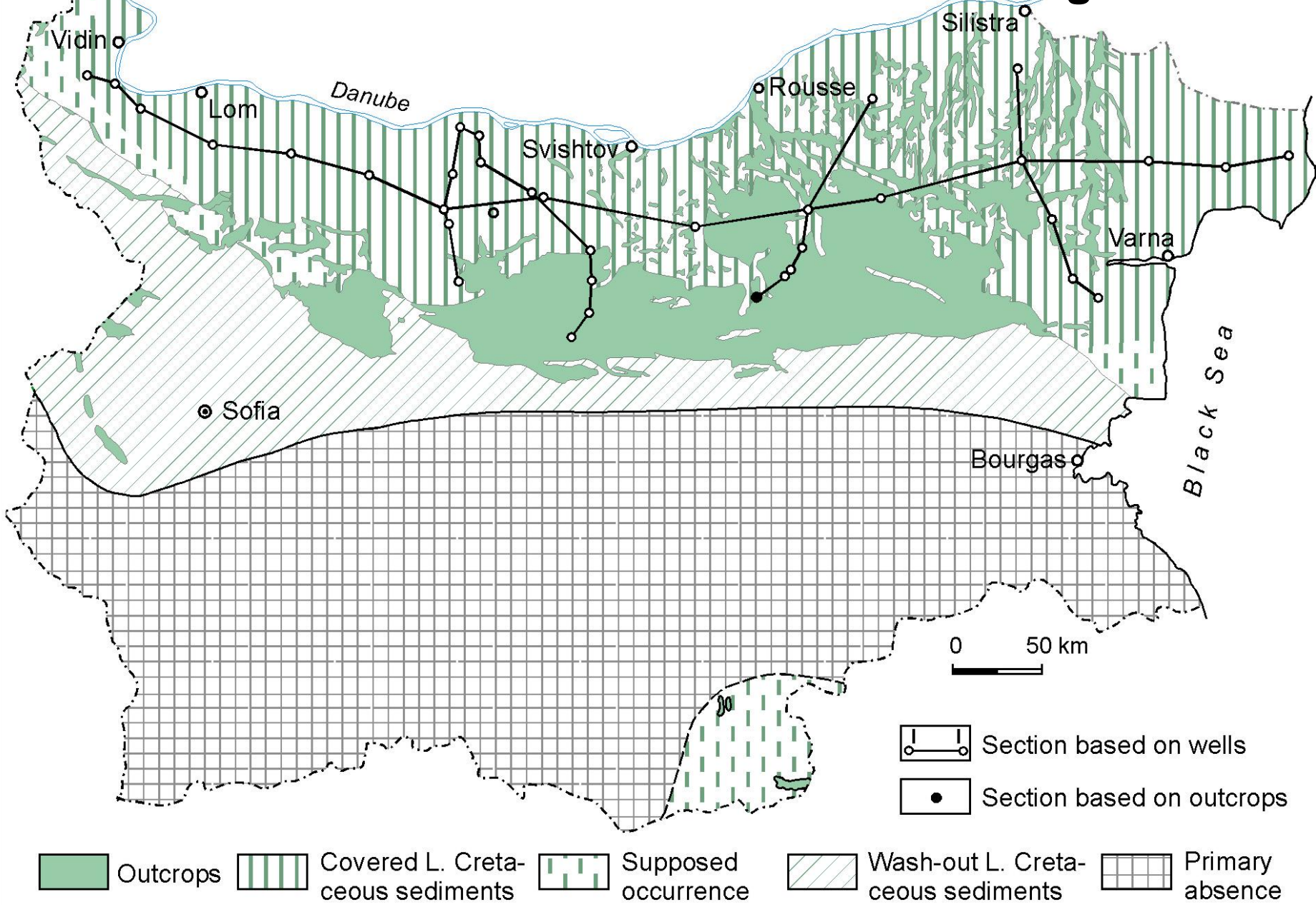
- Dry land (denudation)
- Marine regime without deposition
- Transgressive clastic-pelitic association
- Transgressive terrigenous-limestone association
- Inundation terrigenous-limestone association
- Transgressive terrigenous-radiolarite association
- Section number

0 50 km

End of Tithonian



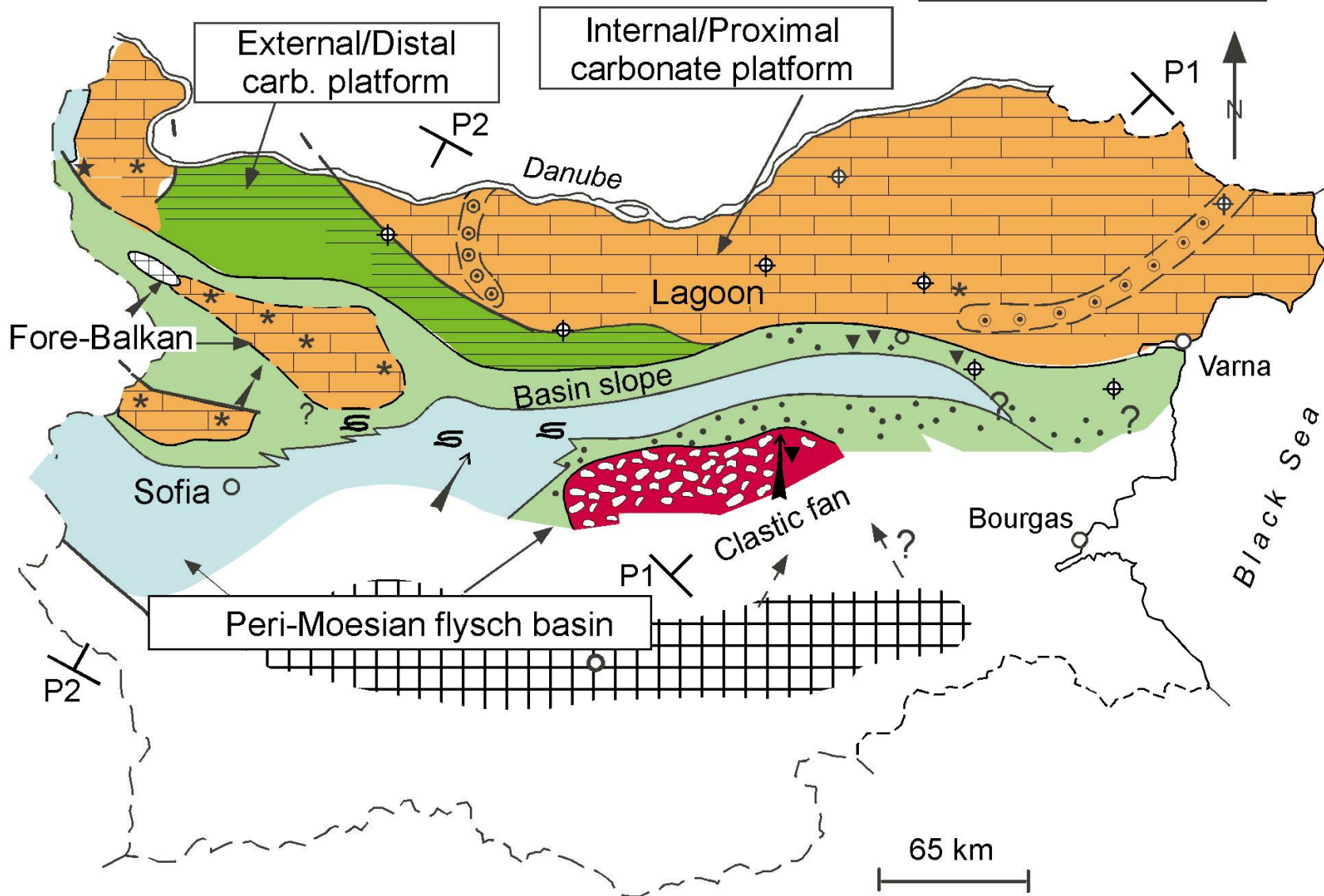
Occurrence of Lower Cretaceous in Bulgaria



Early Cretaceous palaeogeographic maps (without palinospastics)

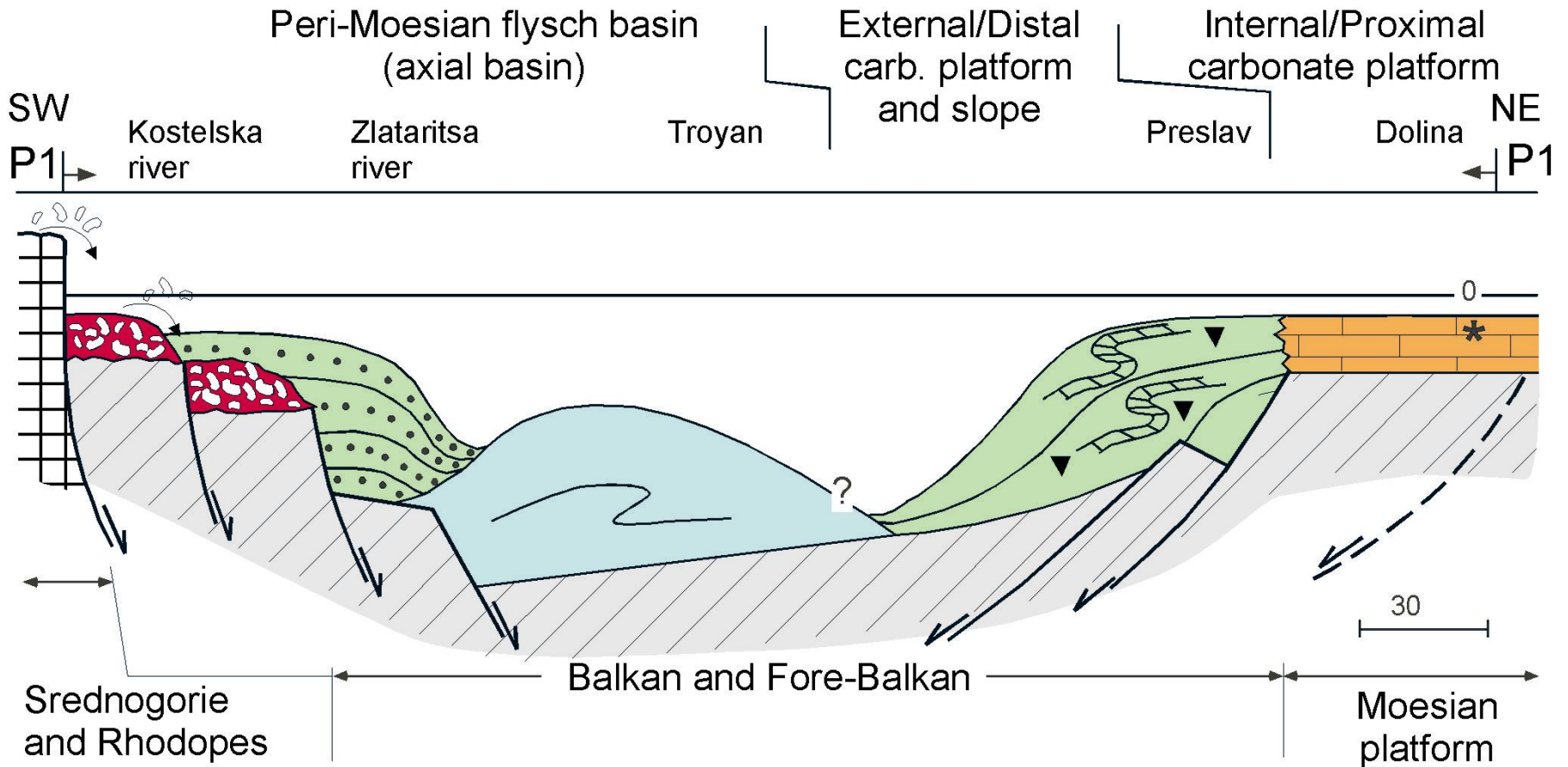
Berriasian

144,2 – 140,2 Ma
ICS – ISChart, 2004



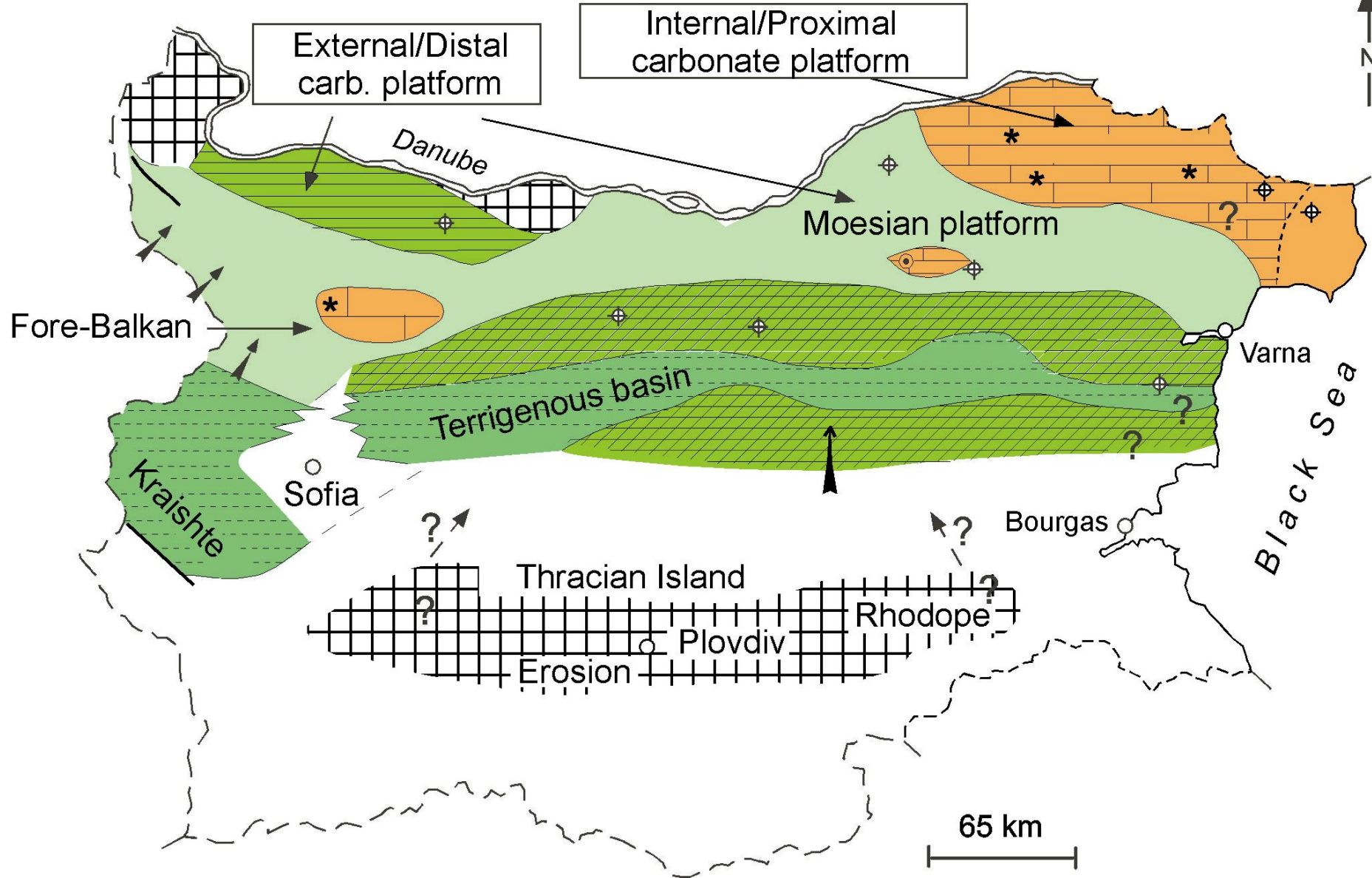
Berriasian P1-P1

Zones *P. grandis* – *T. occitanica*



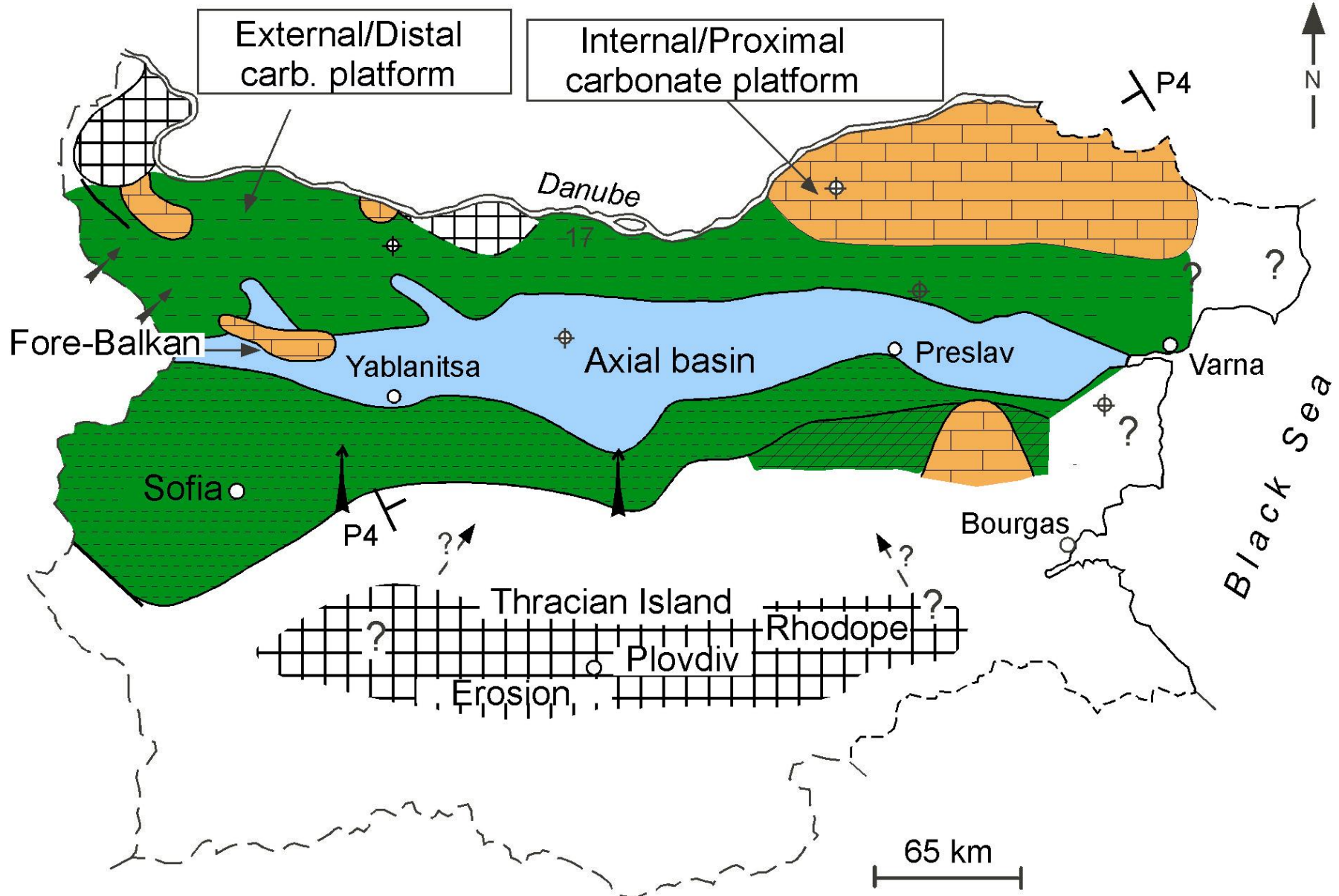
Hauterivian

136,4 – 130 Ma
ICS – ISChart, 2004



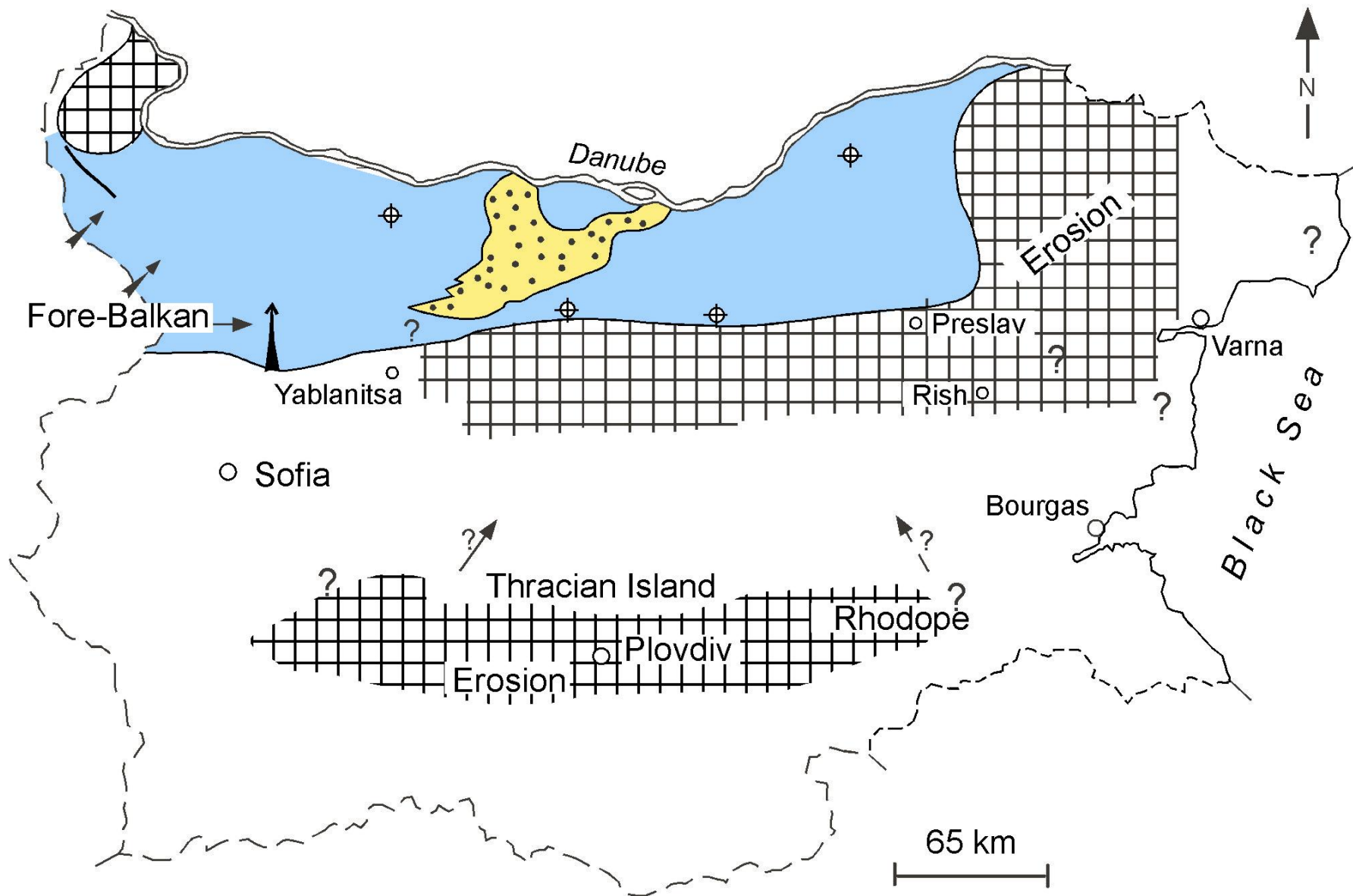
Early Barremian

130 – 128,5 Ma
ICS – ISChart, 2004



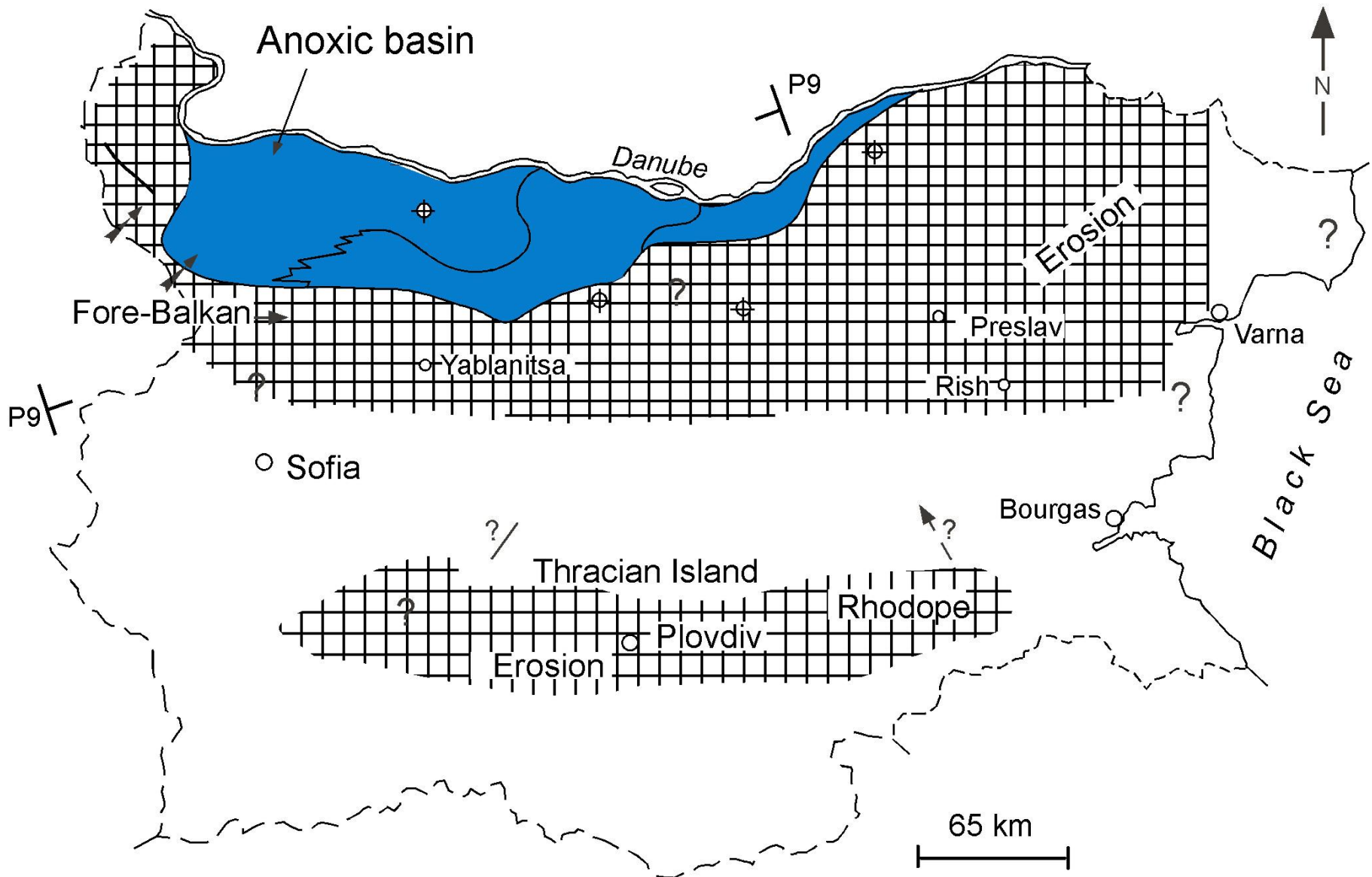
Late Aptian

116 – 112 Ma
ICS – ISChart, 2004

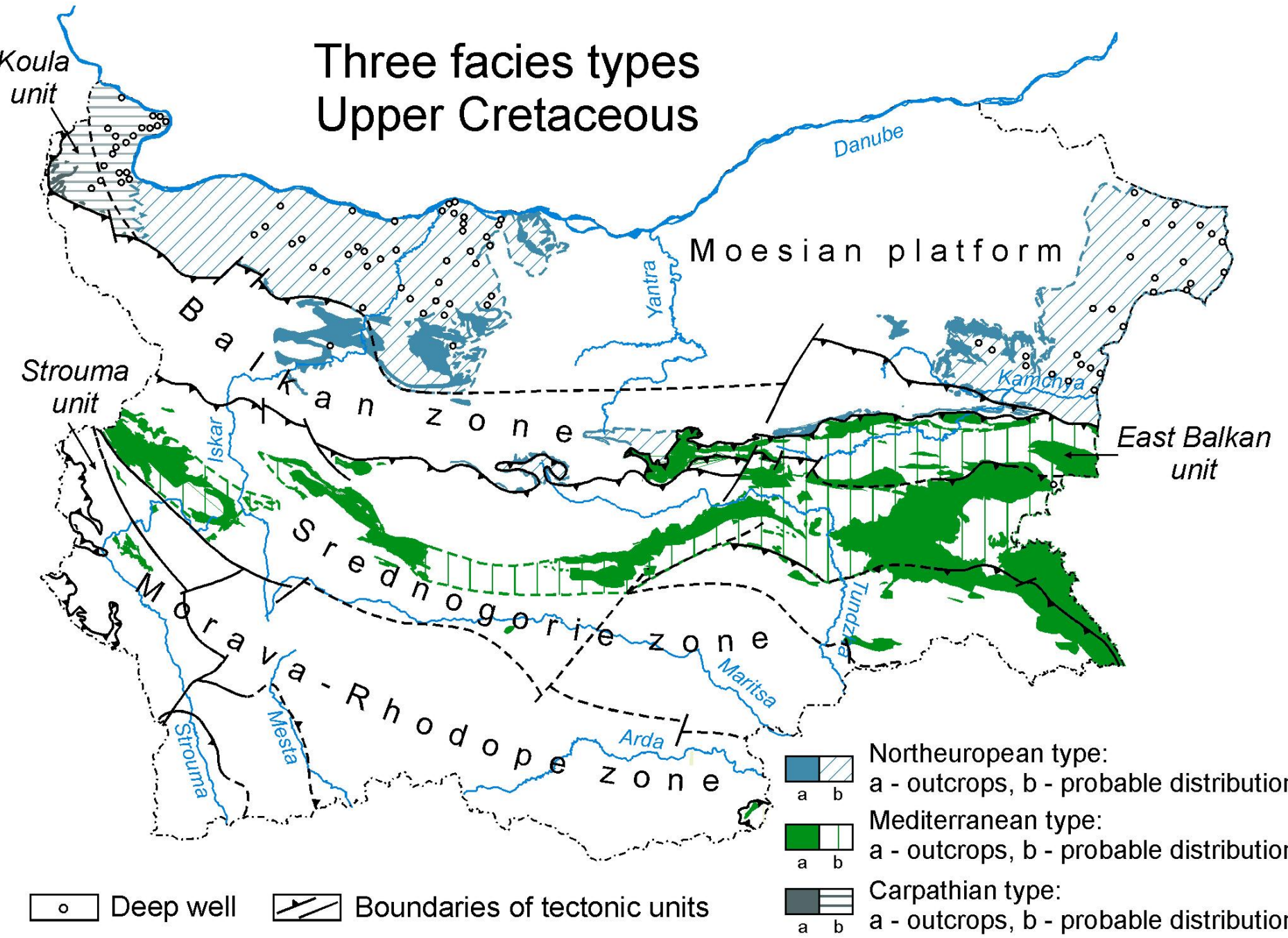


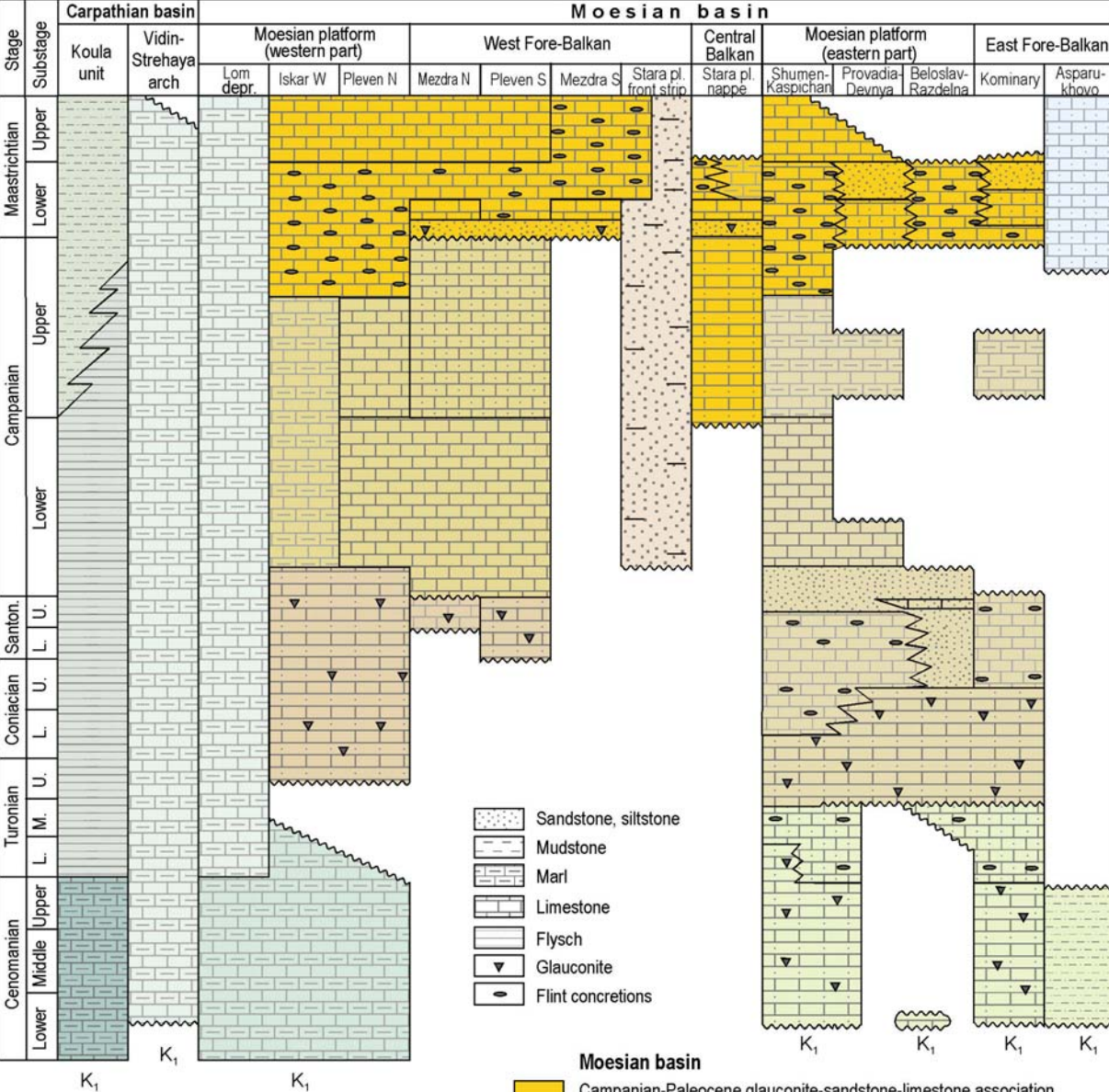
Middle Albian

102,6 – 107 Ma BP
ICS – ISChart, 2004



Three facies types Upper Cretaceous





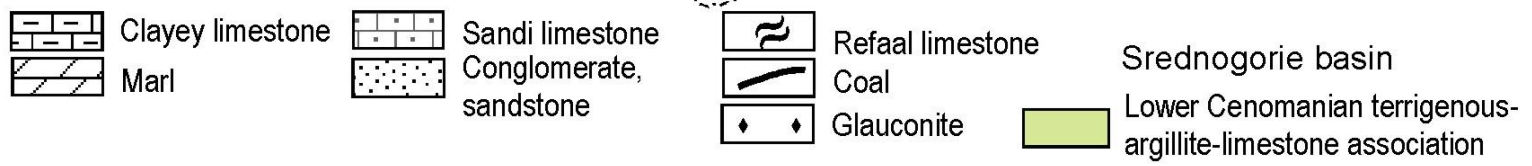
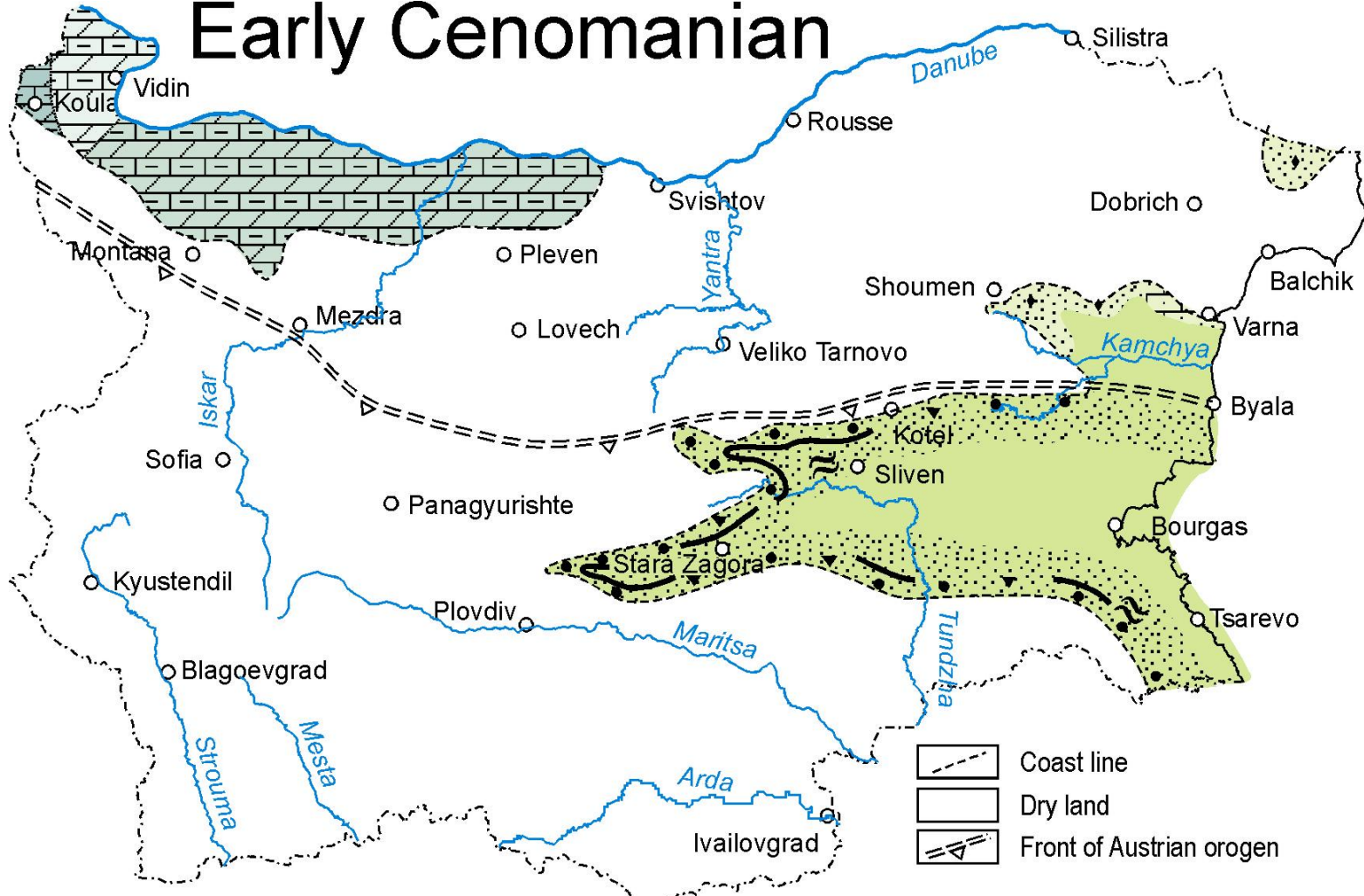
Carpathian basin

- Turonian-Paleocene flysch association
- Cenomanian-Danian limestone-marl association
- Alb-Cenomanian limestone-marl association

Moesian basin

- Campanian-Paleocene glauconite-sandstone-limestone association
- Campanian-Maastrichtian limestone association
- Campanian-Maastrichtian sandstone association
- Campanian terrigenous-limestone association
- Upper Turonian-Campanian sandstone limestone association
- Upper Turonian-Lower Campanian terrigenous-limestone association
- Cenomanian-Turonian sandstone-limestone association
- Cenomanian-Turonian marl association

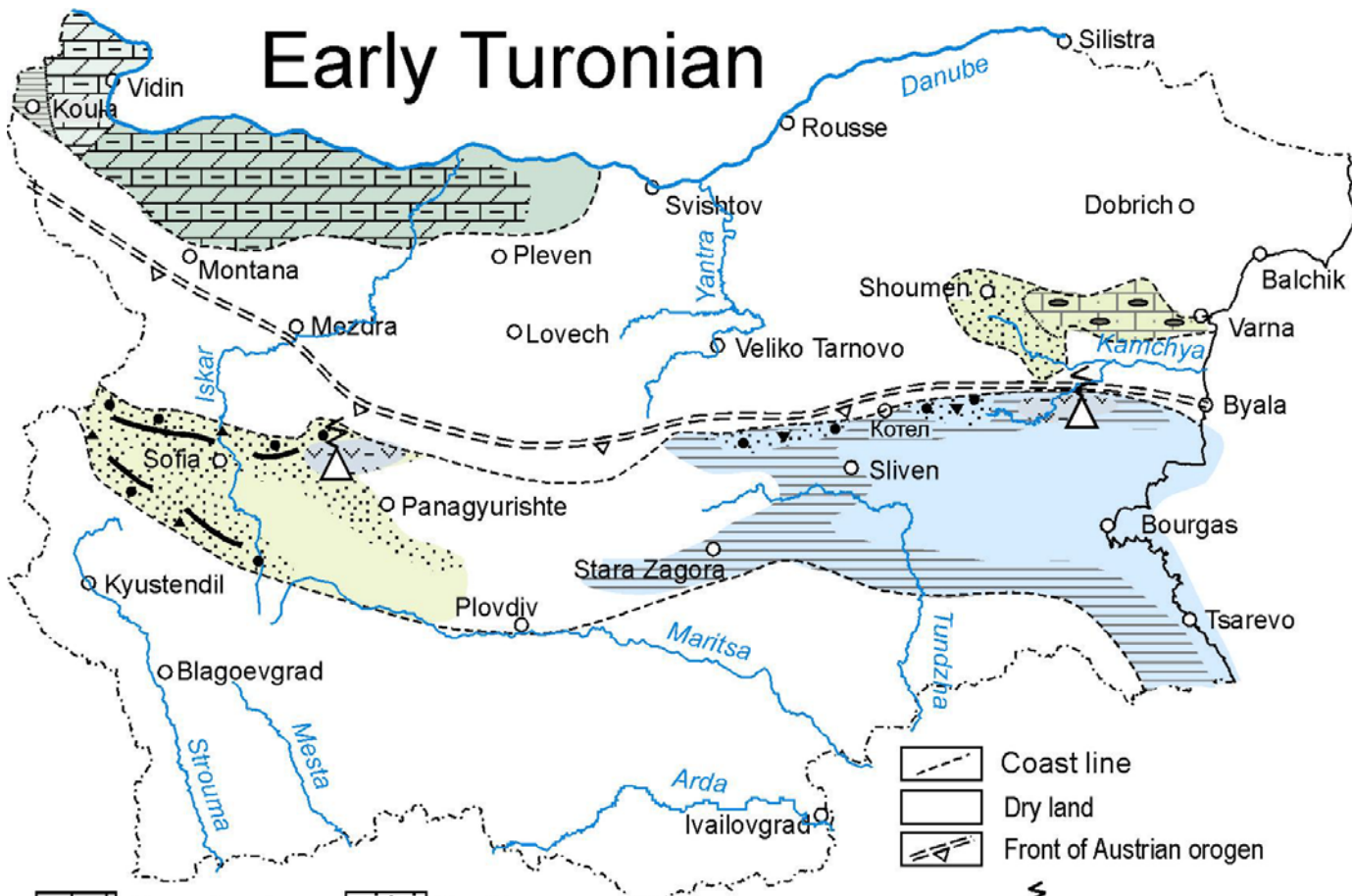
Early Cenomanian



Carpathian and Moesian basin



Early Turonian



- | | | | | | | | |
|--|------------------|--|-----------------------------------|--|-------------------|--|--------------|
| | Clayey limestone | | Sandstone | | Coal | | Paleovolcano |
| | Marl | | Conglomerate, brecciaconglomerate | | Flint concretions | | Volcanics |
| | Sandy limestone | | Glaucconite | | Turbidites | | ± sediments |

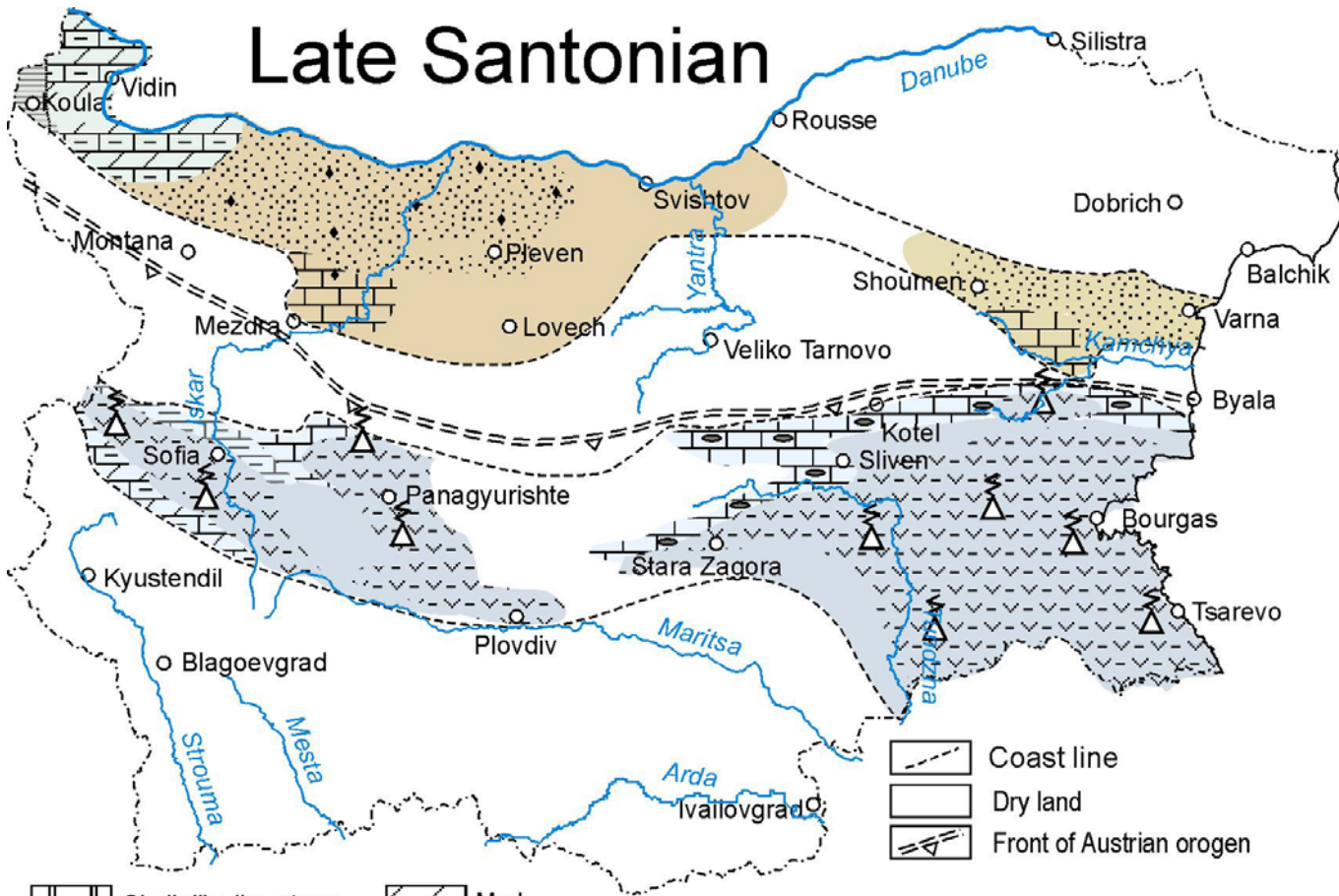
Carpathian and Moesian basin

- | | | | |
|--|---|--|---------------------------------------|
| | Cenomanian-Danian limestone-marl association | | Cenomanian-Turonian marl association |
| | Cenomanian-Turonian sandstone-limestone association | | Turonian-Paleocene flysch association |

Srednogorie basin

- | | | | | | |
|--|--|--|--|--|--|
| | Upper Cenomanian-Campanian volcano-sedimentary association | | Upper Cenomanian-Turonian flysch association | | Turonian terrigenous-limestone association |
|--|--|--|--|--|--|

Late Santonian



- | | | | |
|----------------------|-----------------|--------------------------|-----------------------|
| Chalk-like limestone | Marl | Coast line | Dry land |
| Limestone | Sandy limestone | Front of Austrian orogen | Flint concretions |
| Clayey limestone | Sandstone | Paleovcano | Volcanics ± sediments |
| | Glauconite | | |

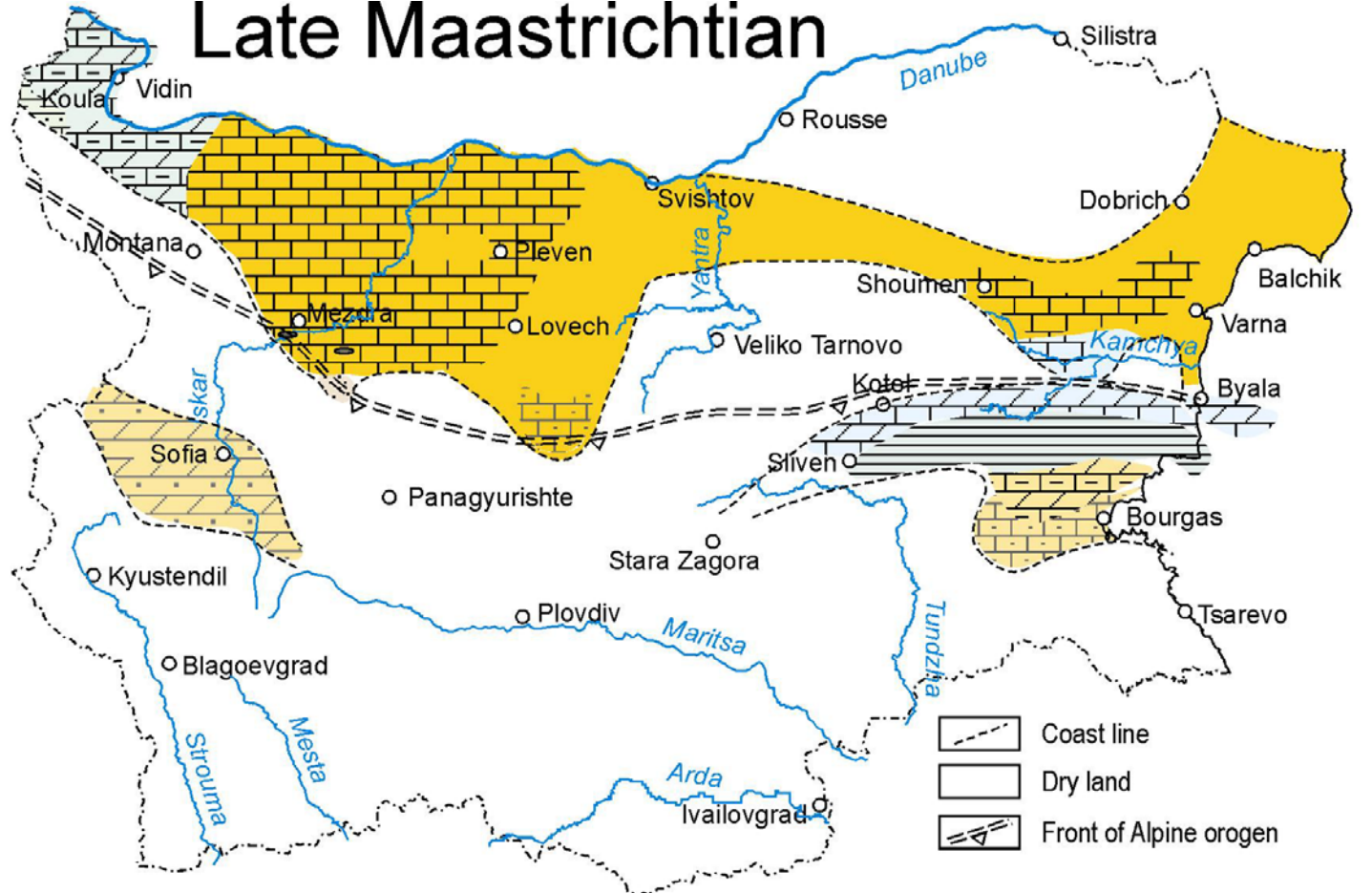
Carpathian and Moesian basin

- | | |
|--|--|
| Cenomanian-Danian limestone-marl association | Upper Turonian-Lower Campanian terrigenous-limestone association |
| Upper Turonian-Campanian sandstone-limestone association | Turonian-Paleocene flysch association |

Srednogorie basin

- | | |
|--|--|
| Upper Cenomanian-Campanian volcano-sedimentary association | Coniacian-Maastrichtian limestone-marl association |
|--|--|

Late Maastrichtian



- | | | | | | |
|--|------------------|--|-----------------|--|-----------------------------------|
| | Turbidites | | Sandy limestone | | Sandstone |
| | Limestone | | Marl | | Conglomerate, brecciaconglomerate |
| | Clayey limestone | | Argillite | | Flint concretions |

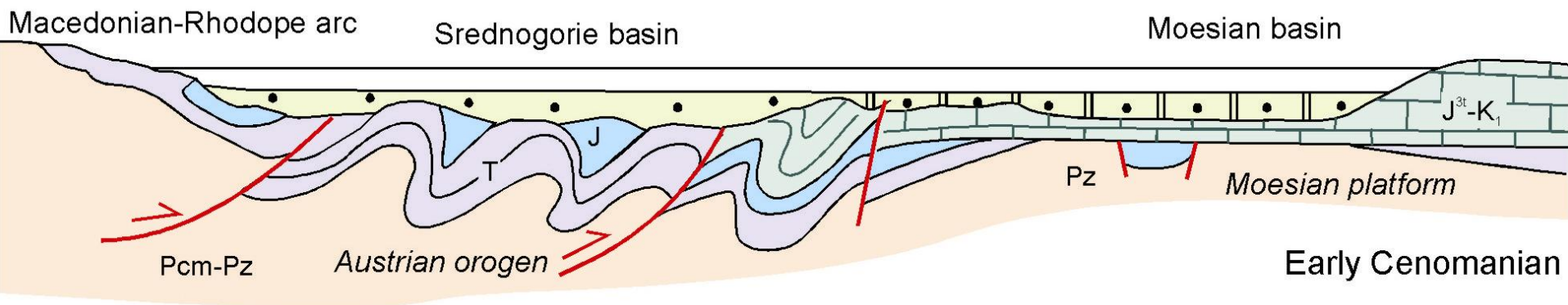
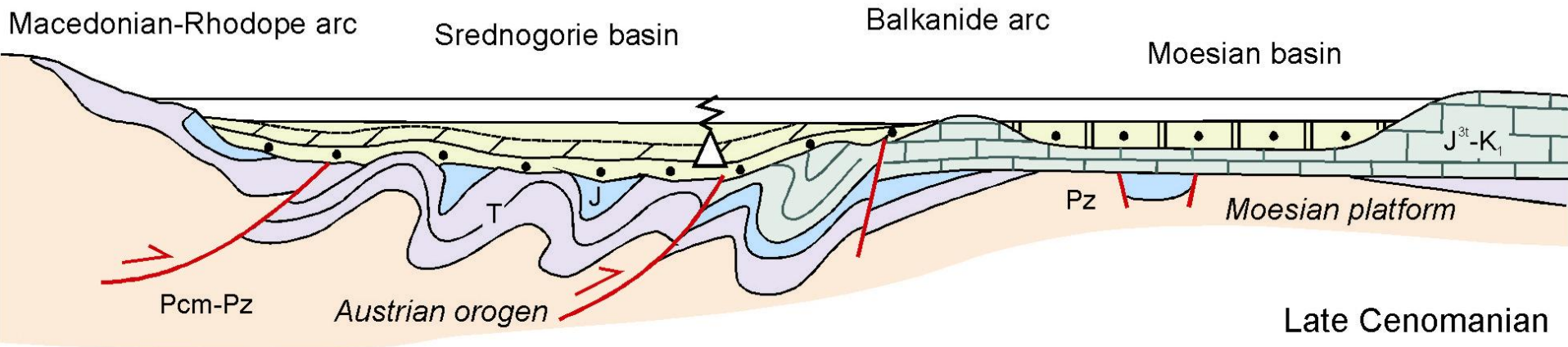
Carpathian and Moesian basin

- Campanian-Paleocene glauconite-sandstone-limestone association
- Turonian-Paleocene flysch association
- Cenomanian-Danian limestone-marl association


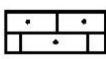





Srednogorie basin

- Maastrichtian terrigenous-limestone association
- Campanian-Maastrichtian flysch association
- Coniacian-Maastrichtian limestone-marl association

Cartoons for the geodynamic evolution in Late Cretaceous times



SREDNOGORIE BASIN

- | | | | |
|---|--|---|---|
|  | U. Cenomanian - Campanian
volcano-sedimentary association |  | Maastrichtian terrigenous-
limestone association |
|  | Upper Cenomanian - Turonian
fysch/flyschoid association |  | Campanian - Maastrichtian
fysch association |
|  | Upper Cenomanian
marl association |  | Coniacian - Maastrichtian
limestone-marl association |
|  | Lower Cenomanian terrigenous-shale-limestone association | | |

MOESIAN BASIN

- | | |
|---|---|
|  | Campanian - Palaeocene glauconite-
sandstone-limestone association |
|  | Upper Turonian - Campanian
sandstone - limestone association |
|  | Cenomanian - Turonian
sandstone - limestone association |

